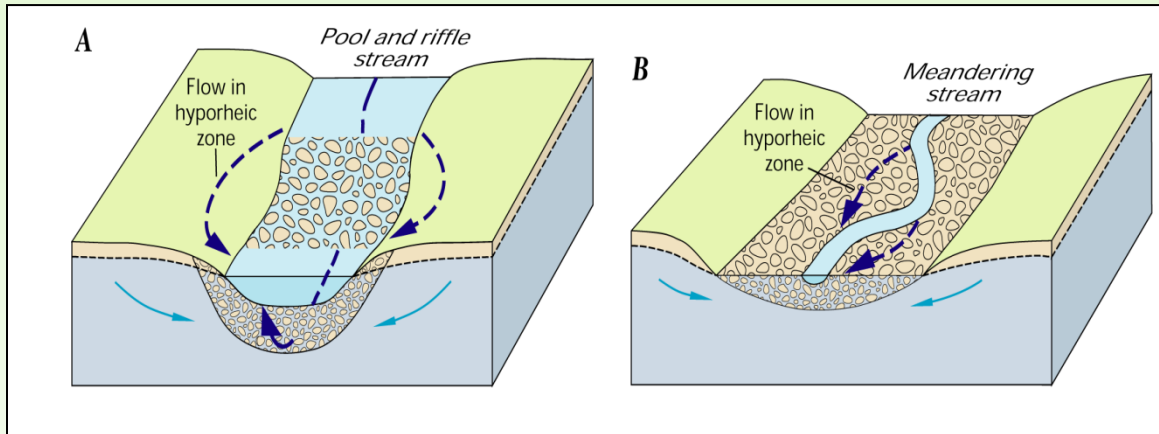


Ground-water surface-water exchange in fluvial settings

Take-home message: Be constantly aware of scale

Hyporheic exchange

Masaki Hayashi,
hayashi@ucalgary.ca
Donald Rosenberry,
rosenberry@mines.edu



Definition of hyporheic exchange

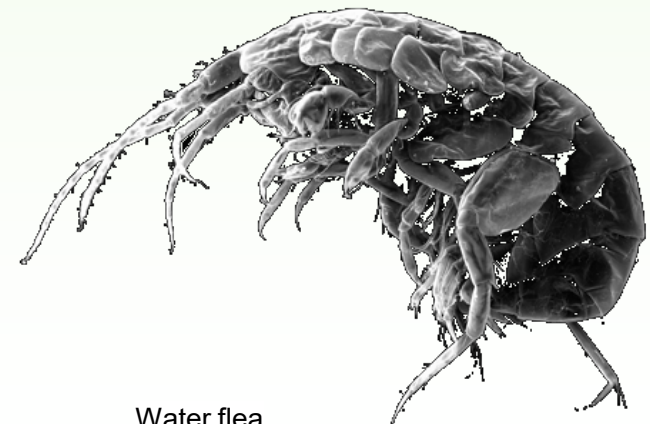
“The region of saturated sediments beneath and beside the active channel and that contains some proportion of surface water that was part of the flow in the surface channel and went back underground and can mix with groundwater” – California Dept. of Water Resources GW glossary

“The subsurface zone where stream water flows through short segments of its adjacent bed and banks is referred to as the hyporheic zone.” – USGS Circular 1139

Consider the scale of the process relative to the scale of the measurement.

There are many definitions – these are a couple that we like.

Important ecotone



Water flea

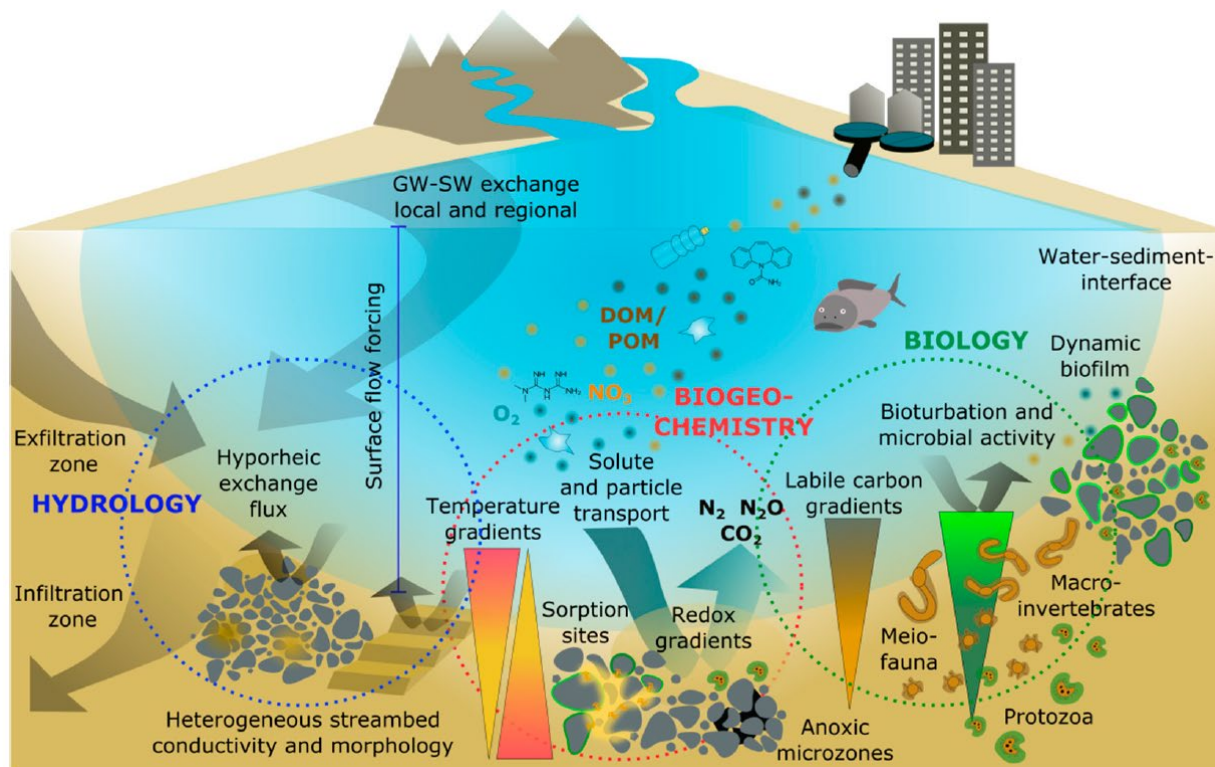
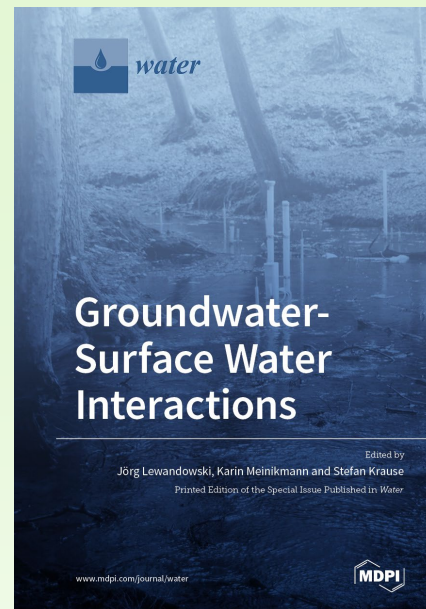
1 mm

Is the Hyporheic Zone Relevant beyond the Scientific Community?

Jörg Lewandowski ^{1,2,*}, Shai Arnon ³, Eddie Banks ⁴, Okke Batelaan ⁴, Andrea Betterle ^{5,6}, Tabea Broecker ⁷, Claudia Coll ⁸, Jennifer D. Drummond ⁹, Jaime Gaona Garcia ^{1,10,11}, Jason Galloway ^{1,2}, Jesus Gomez-Velez ¹², Robert C. Grabowski ¹³, Skuyler P. Herzog ¹⁴, Reinhard Hinkelmann ⁷, Anja Höhne ^{1,15}, Juliane Hollender ⁵, Marcus A. Horn ^{16,17}, Anna Jaeger ^{1,2}, Stefan Krause ⁹, Adrian Löchner Prats ¹⁸, Chiara Magliozzi ^{13,19}, Karin Meinikmann ^{1,20}, Brian Babak Mojarrad ²¹, Birgit Maria Mueller ^{1,22}, Ignacio Peralta-Maraver ²³, Andrea L. Popp ^{5,24}, Malte Posselt ⁸, Anke Putschew ²², Michael Radke ²⁵, Muhammad Raza ^{26,27}, Joakim Riml ²¹, Anne Robertson ²³, Cyrus Rutere ¹⁶, Jonas L. Schaper ^{1,22}, Mario Schirmer ⁵, Hanna Schulz ^{1,2}, Margaret Shanafield ⁴, Tanu Singh ⁹, Adam S. Ward ¹⁴, Philipp Wolke ^{1,28}, Anders Wörman ²¹ and Liwen Wu ^{1,2}

Papers in this special issue of the journal *Water* on GW-SW exchange give a fairly thorough and recent overview of hyporheic exchange and the broad range of processes that are involved and/or affected. This special issue is a nice resource that summarizes recent findings on this topic.

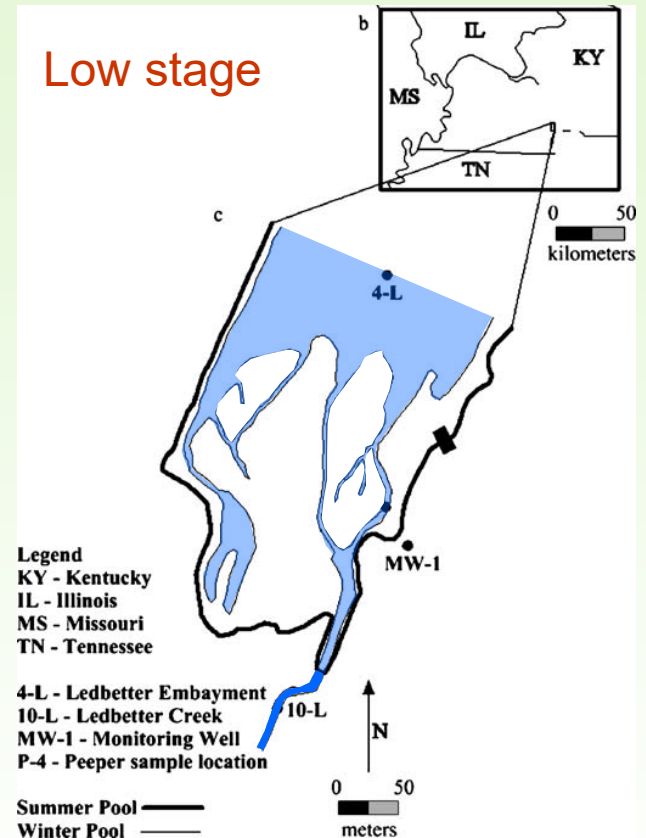
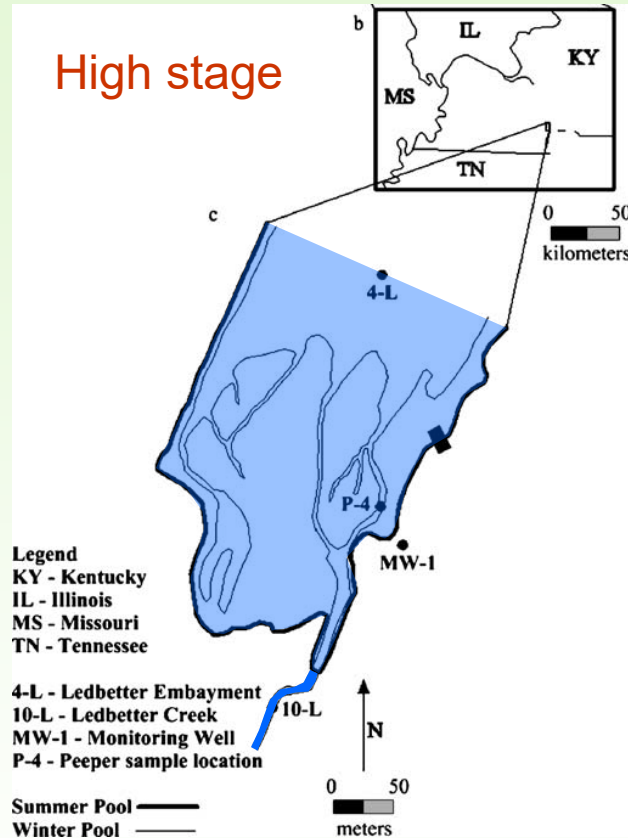
Lewandowski et al.,
2020, *Water*



Hypolentic zone?

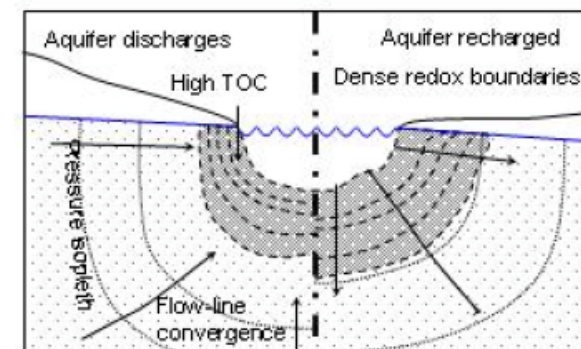
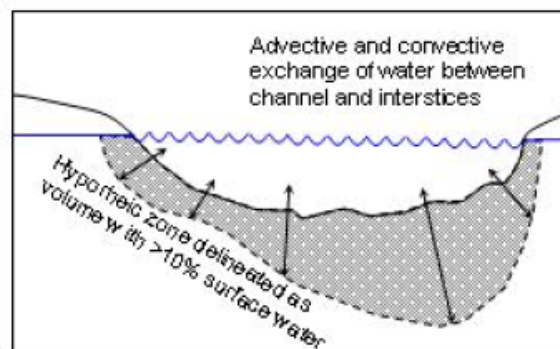
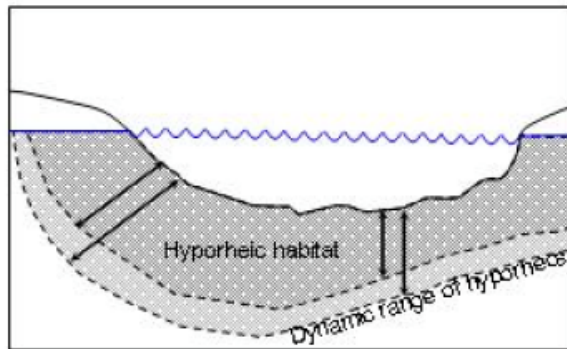
These types of exchanges also occur in lakes and wetlands, but usually to a lesser extent. In the case of lakes, it was first called hypolentic exchange by Thomas Winter. This subsequent paper used water isotopes to distinguish hypolentic water from hyporheic water, where a river flowed into a reservoir. Hyporheic exchange was created by streambed topography, but hypolentic exchange was created by varying reservoir stage.

Where does hyporheic exchange transition to hypolentic exchange?

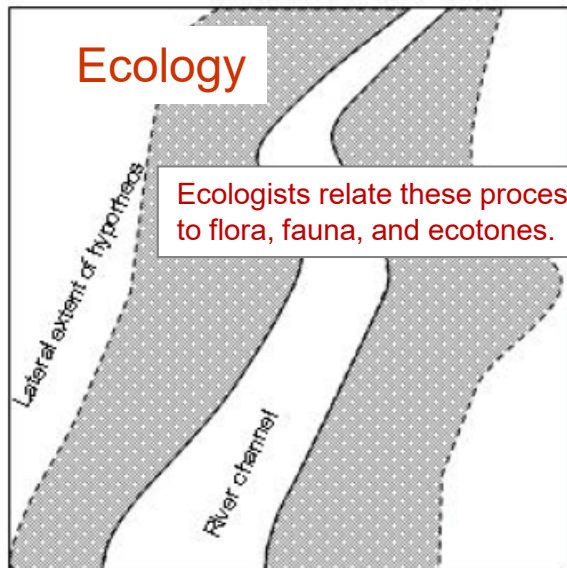


Definition depends on your discipline perspective and the scale of interest

Cross-section

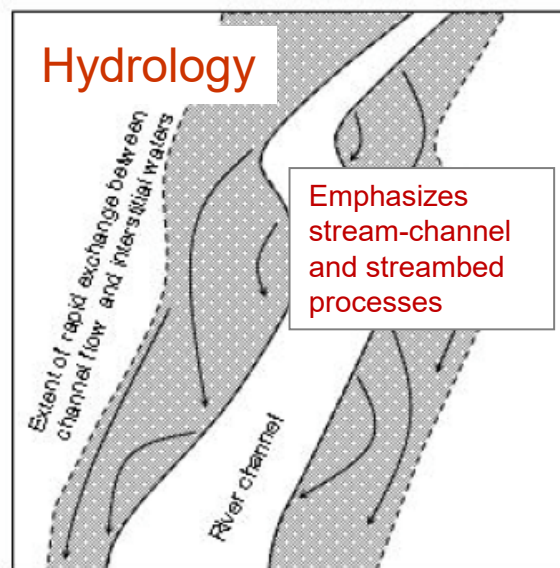


Plan view



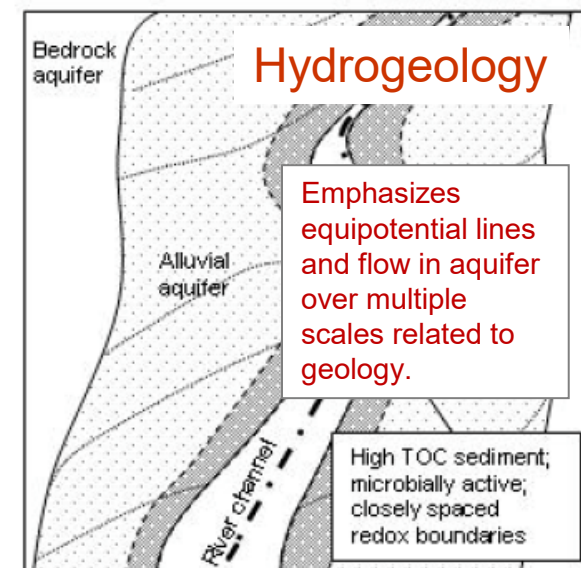
Ecology

Ecologists relate these processes to flora, fauna, and ecotones.



Hydrology

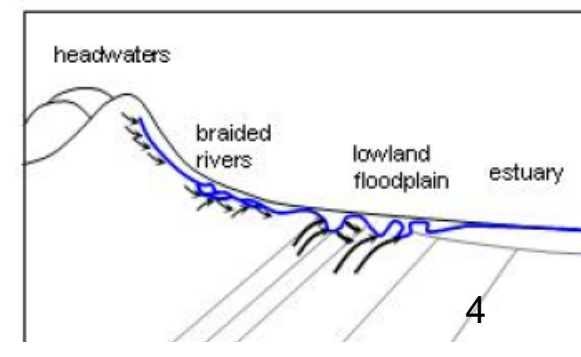
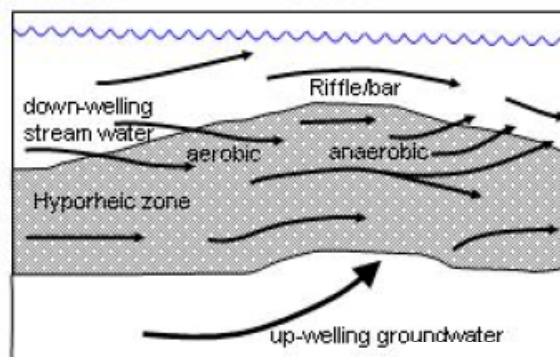
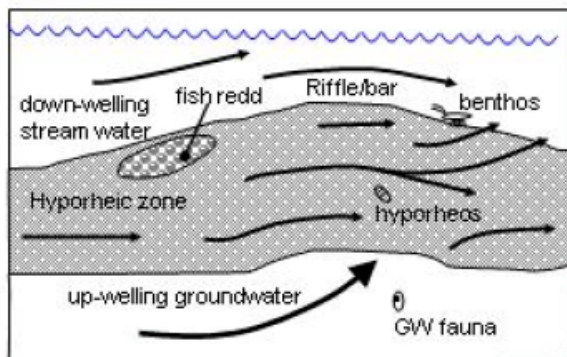
Emphasizes stream-channel and streambed processes



Hydrogeology

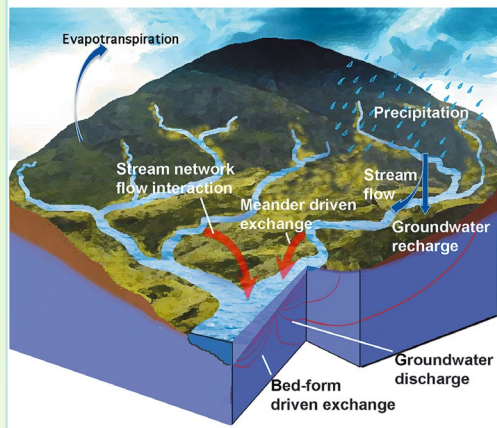
Emphasizes equipotential lines and flow in aquifer over multiple scales related to geology.

Longitudinal section



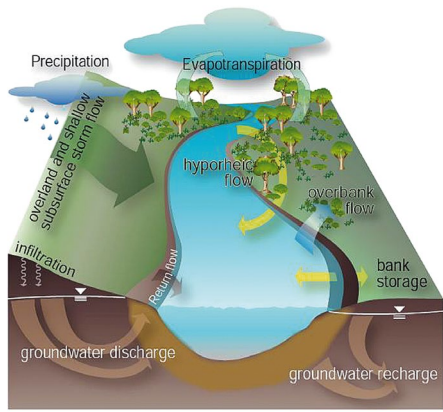
Hyporheic exchange occurs at multiple scales

a) Watershed or Basin scale

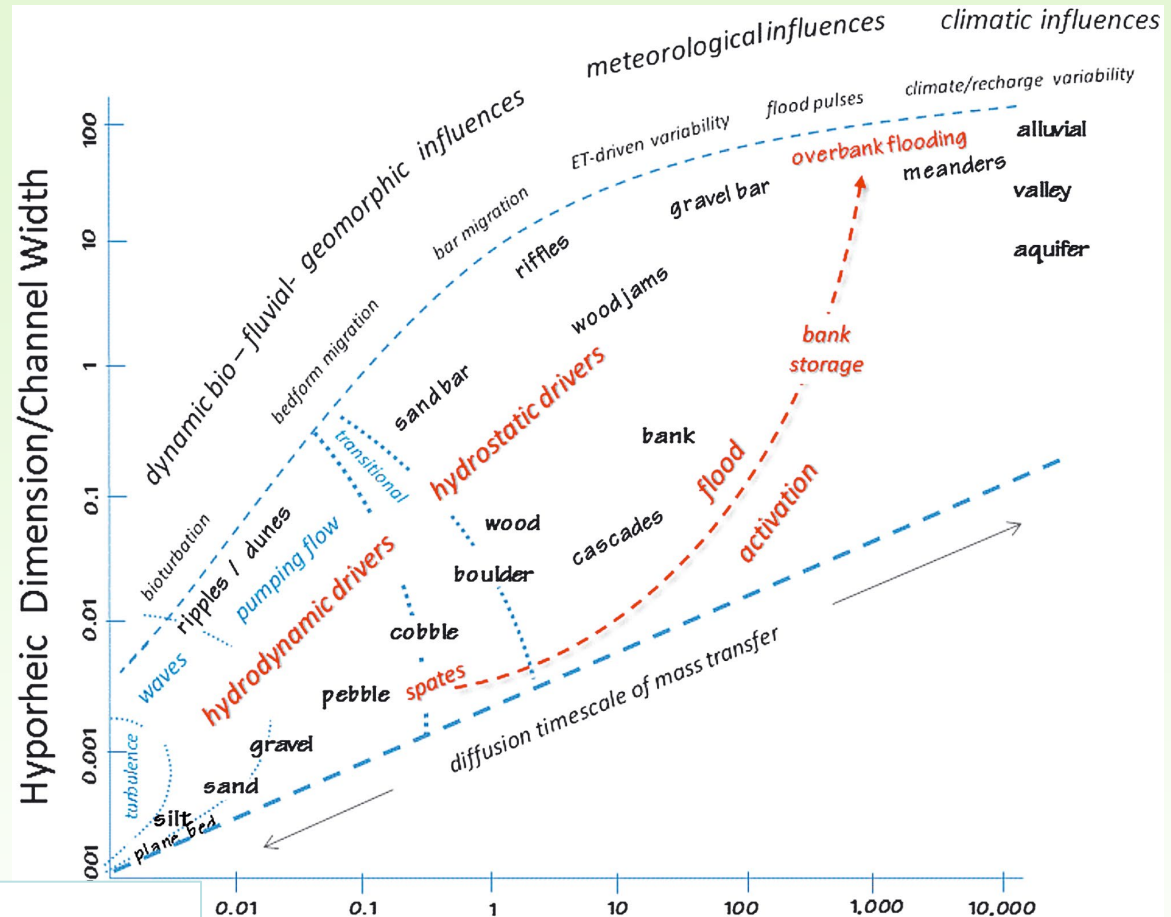
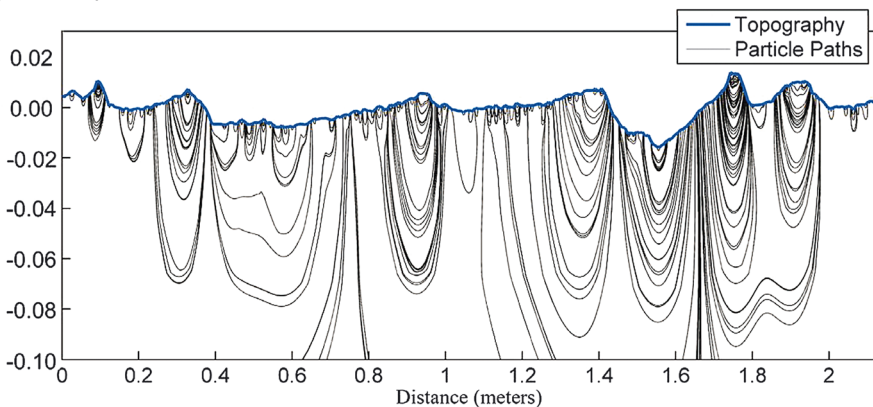


This scale-time plot provides a nice overview of the scale of many of the processes or events that are relevant to hyporheic exchange.

b) River Corridor scale



c) Geomorphic Unit scale



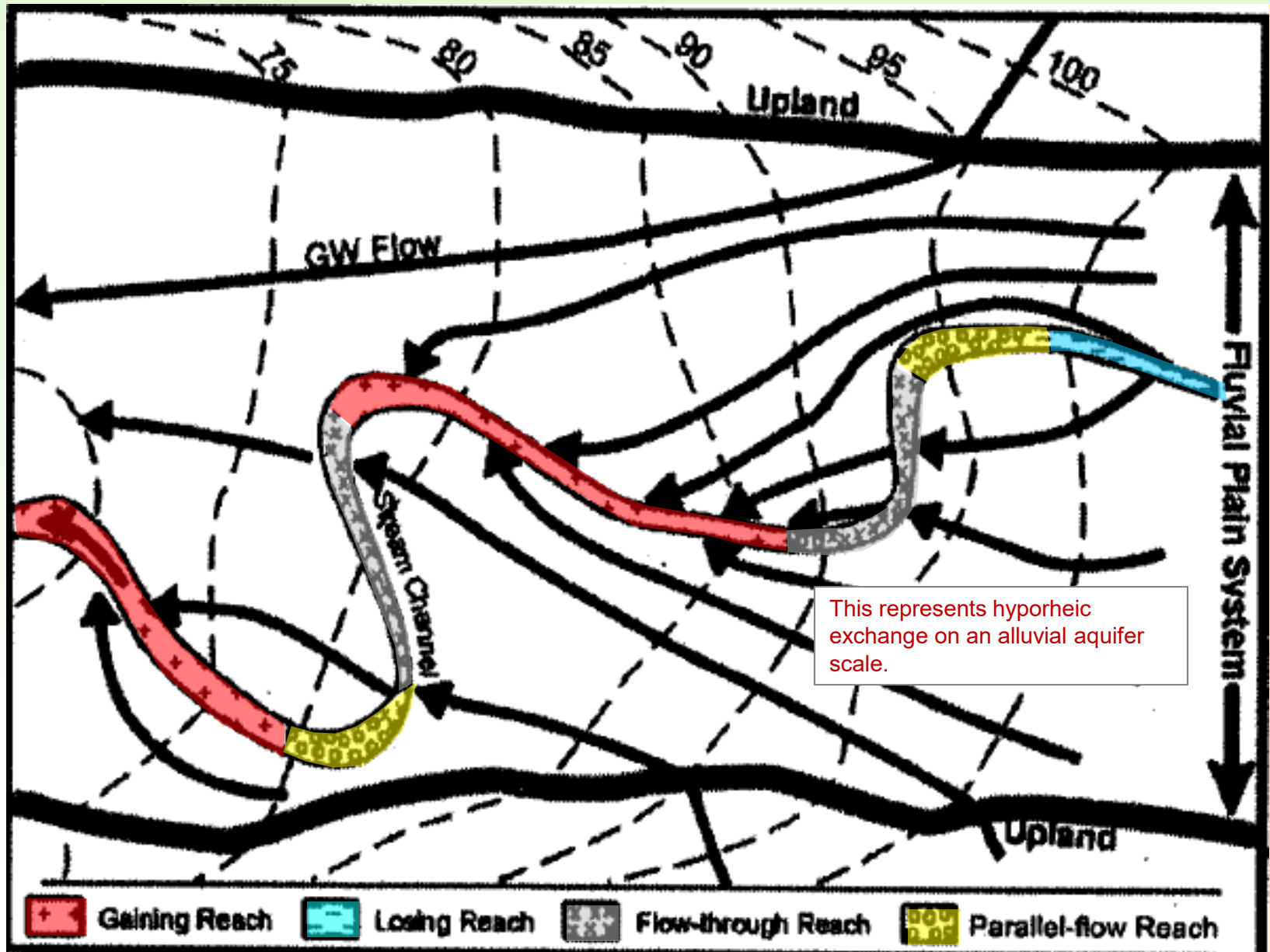
Residence Time of Hyporheic Flow (hour)

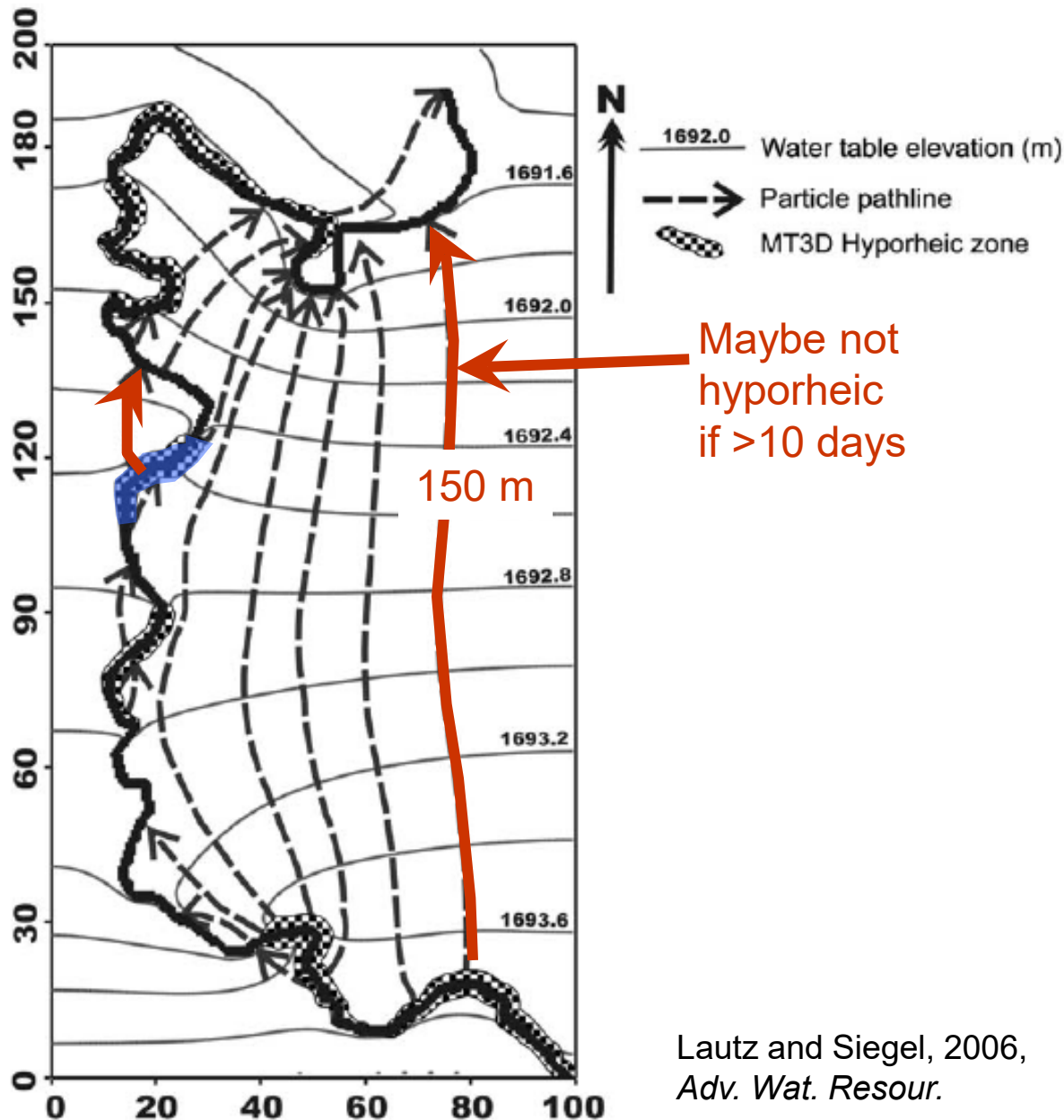
- Basin scale
- Bar scale, meander scale
- Pool-riffle scale
- Dune, bedform scale
- **Let's look at a few examples.**

Meander or floodplain scale involves

1. Parallel flow
2. Flow through
3. Gaining reaches
4. Losing reaches

Woessner, 2000,
Ground Water





Lautz and Siegel, 2006,
Adv. Wat. Resour.

Exchange depends on
perspective and
definition

**Larger-scale exchange
based on MODFLOW
calibrated to wells**

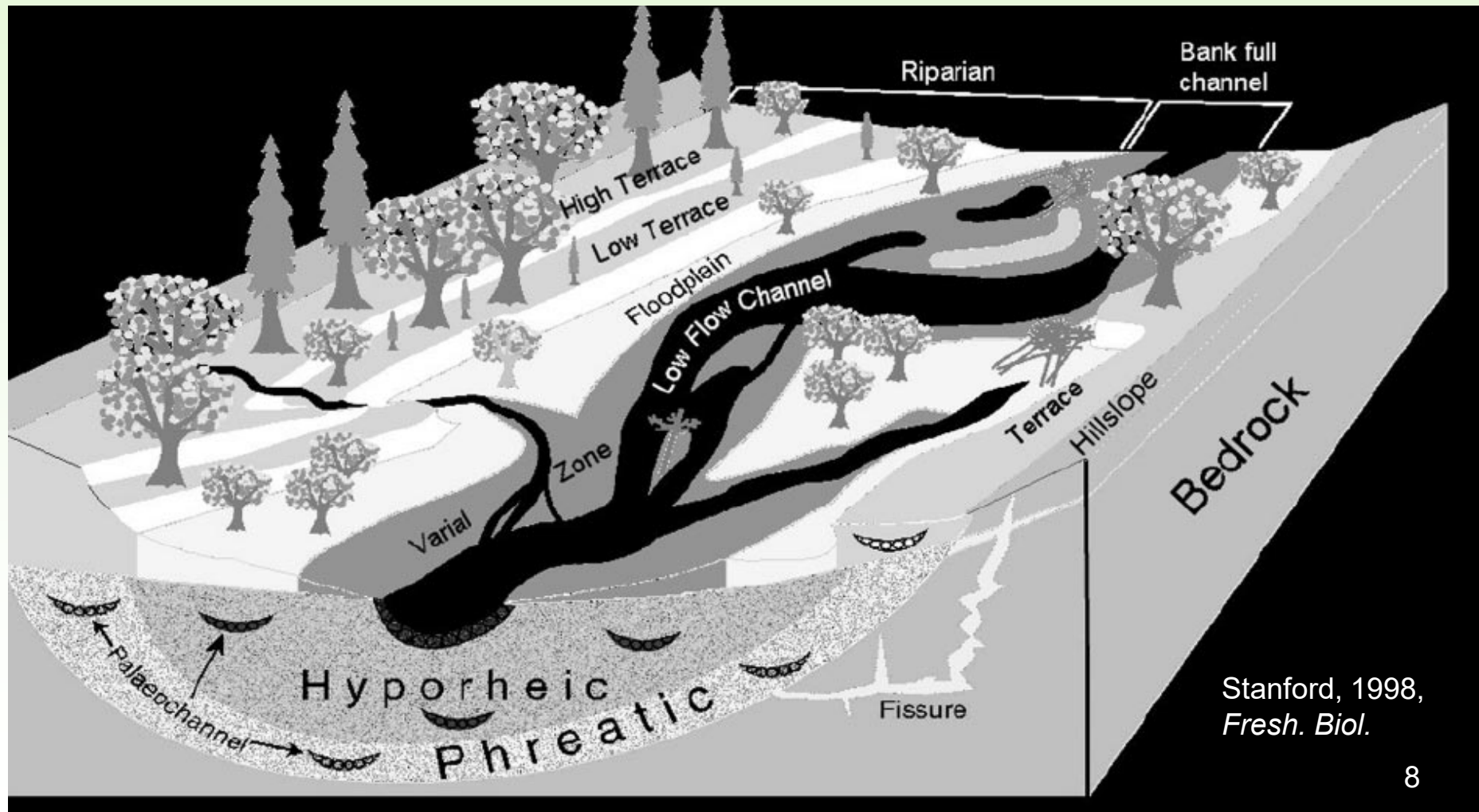
**Local-scale exchange based
on new definition (10%
surface water within a
10-day travel time**

We keep developing new definitions that depend on the perspective of the scientist, this one based on hydrology and geochemistry. The authors used MODFLOW to determine hyporheic zone based on travel times.

Hyporheic zone based on hyporheos

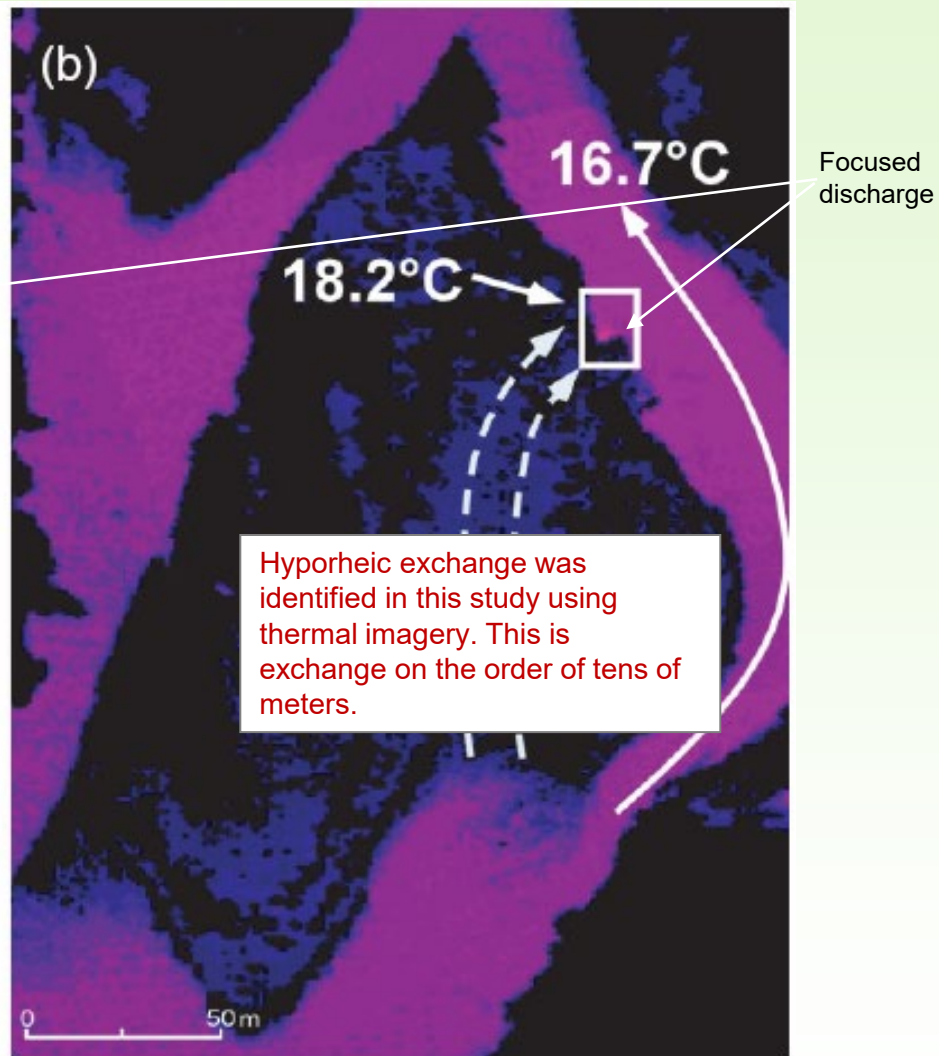
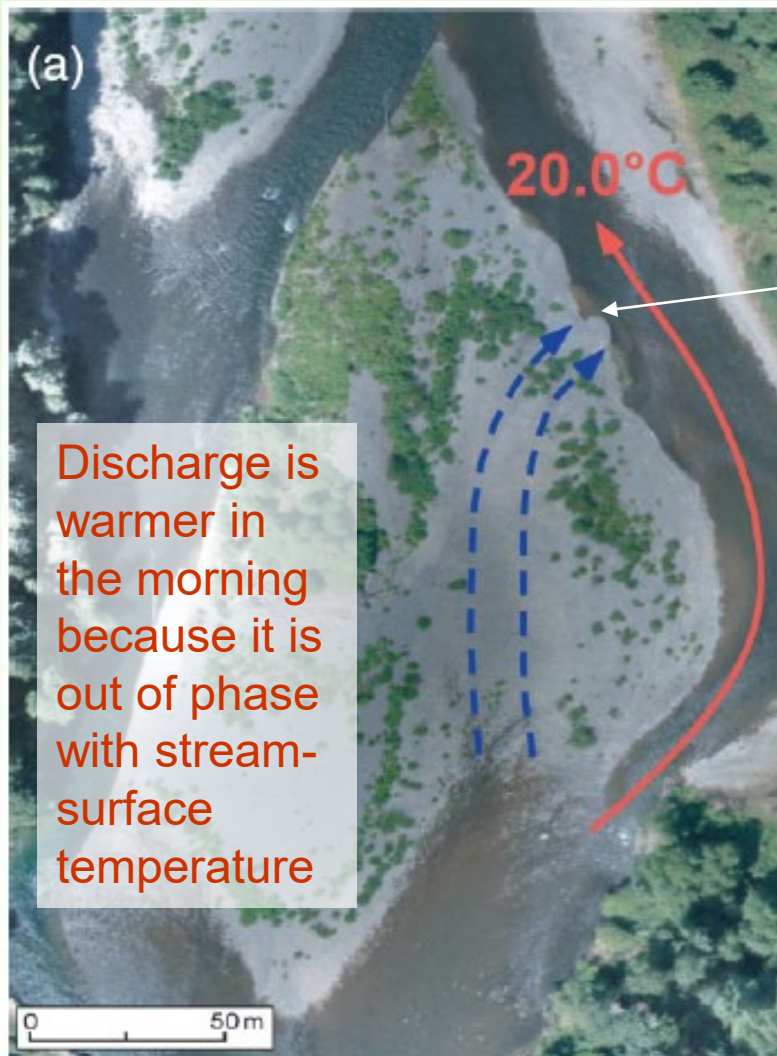
Benthic invertebrates can be found several km away from thalweg

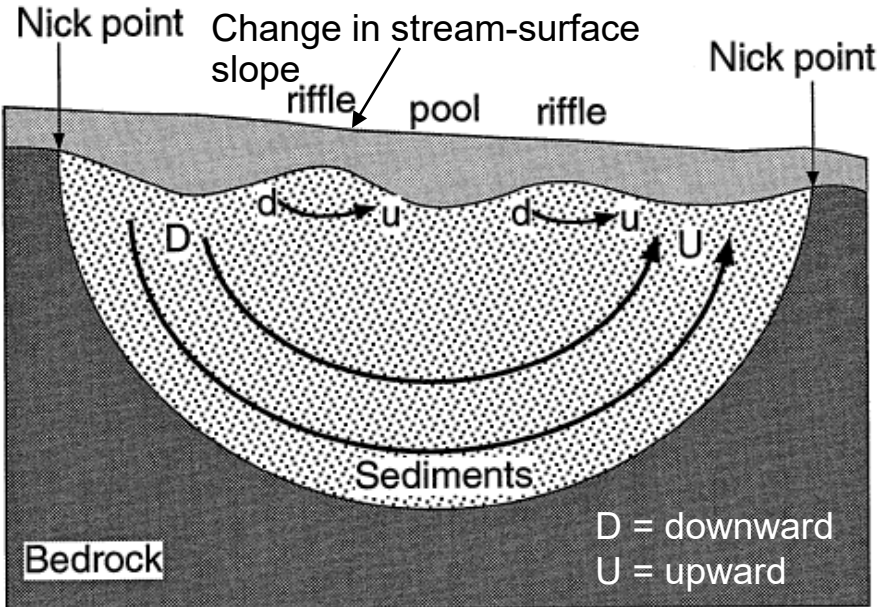
This definition is based on the presence of animals in alluvial sediments that normally are associated with streambed sediments. Sometimes they can be found hundreds of meters away from the active stream channel. Aren't you glad the scale distance in the figure to the right is 1 mm and not 1 m?



Stanford, 1998,
Fresh. Biol.

Bar-scale hyporheic exchange





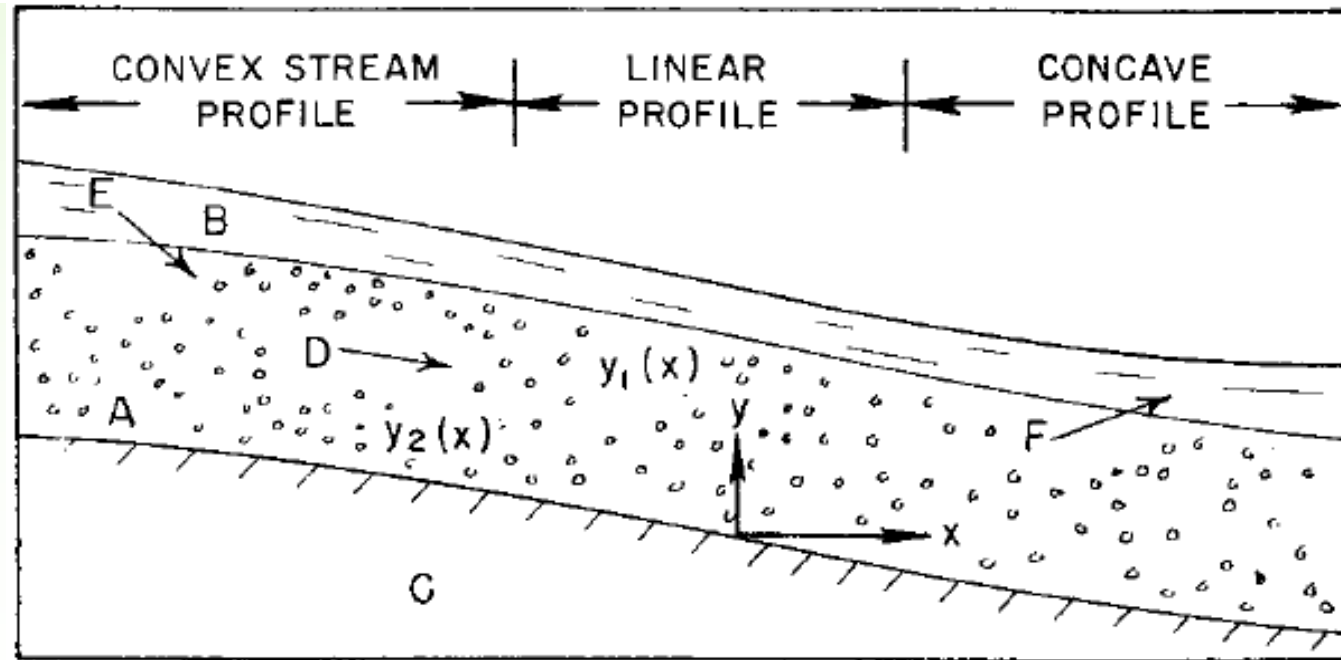
Pool-riffle scale

- Downward seepage at downstream ends of pools
- Upward seepage at upstream ends of pools
- A mix of downward, upward, and bed-parallel seepage in riffles

Brunke and Gonser, 1997,
Fresh. Biol.

The concept of the downstream streambed profile affecting hyporheic exchange has been around for a long time

They didn't call it hyporheic exchange when this was published in 1968.



Vaux, 1968, *Fishery Bulletin*

This is an early publication that illustrates the hyporheic zone.

Nick point: an abrupt change in slope

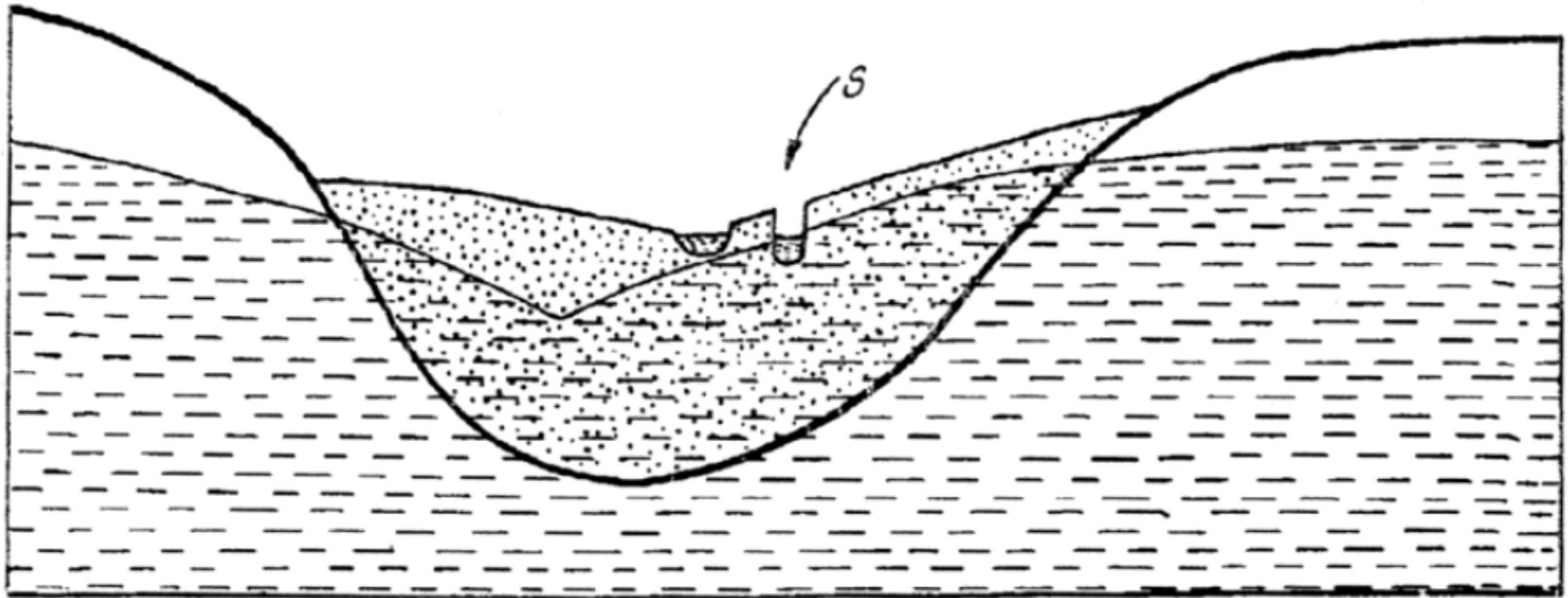
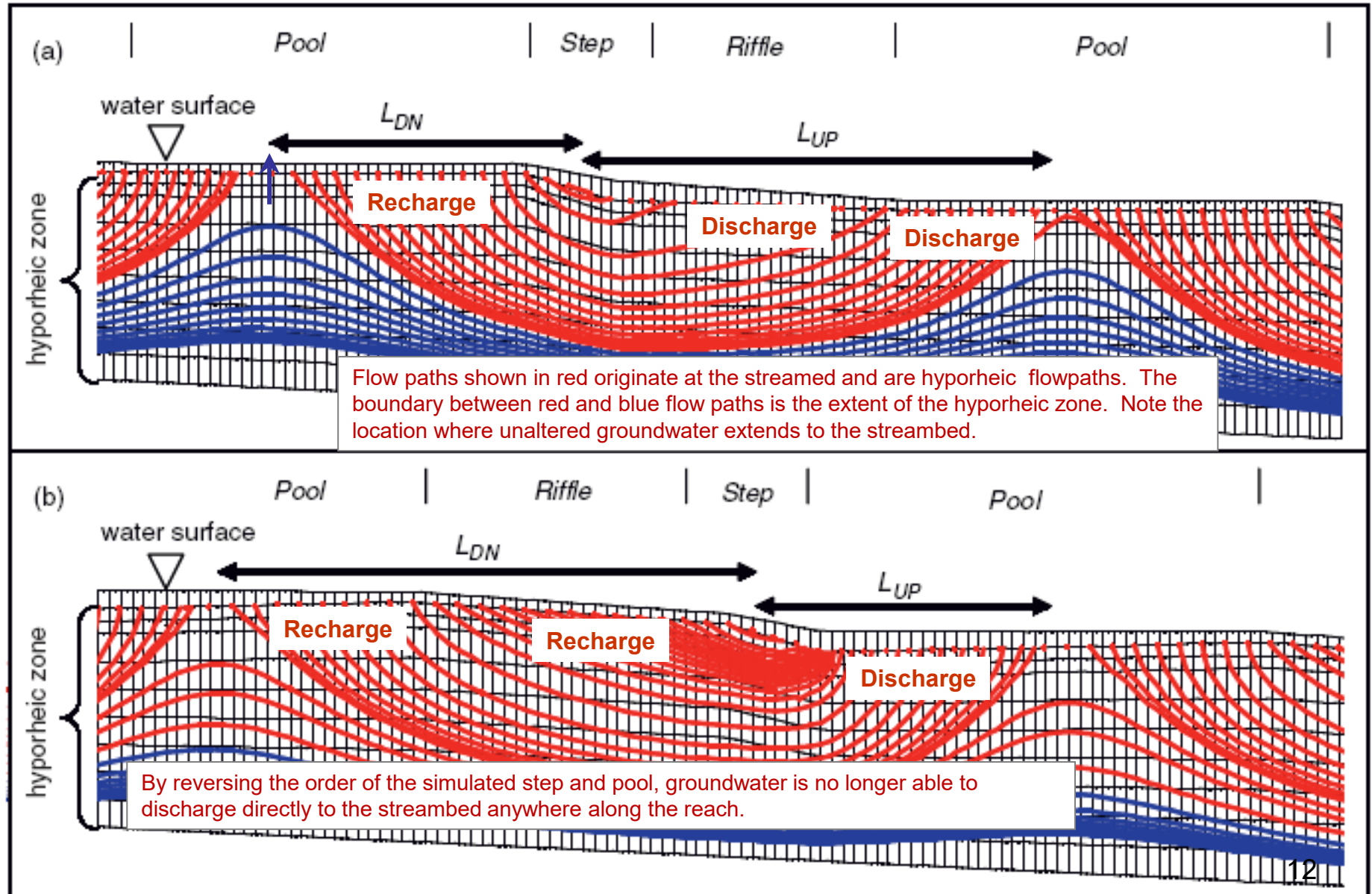


Figure 1. Permeable valley floor; hyporheic water comes both from the stream and the phreatic groundwater (modified after Imbeaux). S, excavation site.

But the earliest mention of the hyporheic zone was by Orghidan in 1959, translated from German to English by Daniel Kaser, 2010.

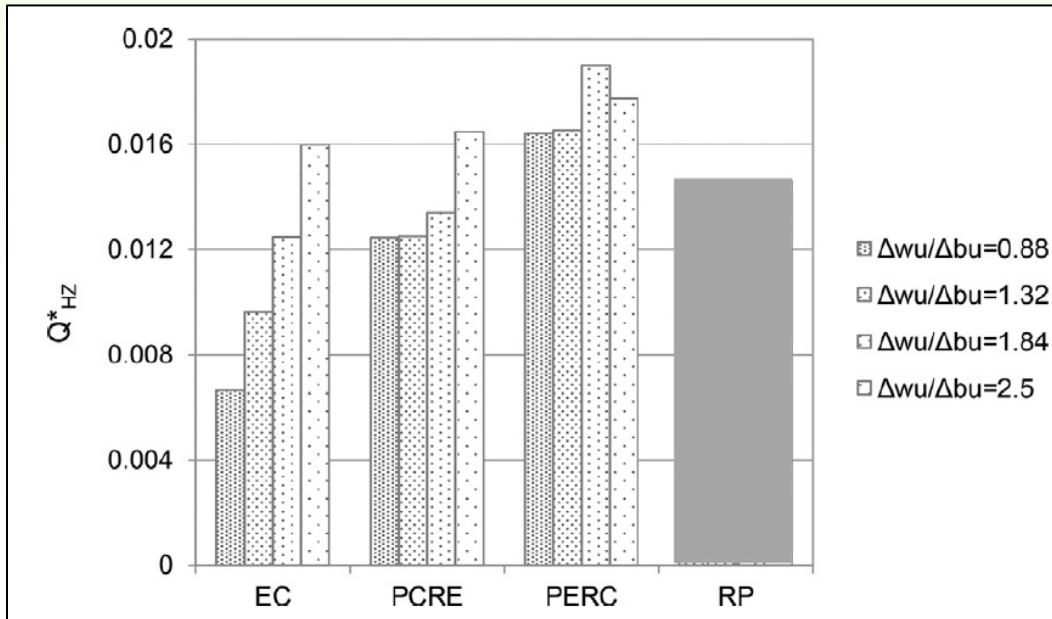
Direction of exchange depends on the degree and sequence of changes in bed slope



Changes in stream width can also induce hyporheic exchange



Variations in stream width can also induce hyporheic exchange, although the influence is generally less than changes in bed slope. The greatest hyporheic exchange occurs when the pool is wide and the riffle is narrow (PERC).



wu = width undulation

bu = bed slope undulation

EC = expansion-constriction

PCRE = pool constricted,
riffle expanded

PERC = pool expanded, riffle
constricted

RP = riffle-pool

Movahedi et al., 2021, *Advances in Water Resources*

Quantification of hyporheic exchange and determination of flowpaths can require intensive instrumentation in and adjacent to the stream

There now are many studies on a scale of cm to m in the literature.



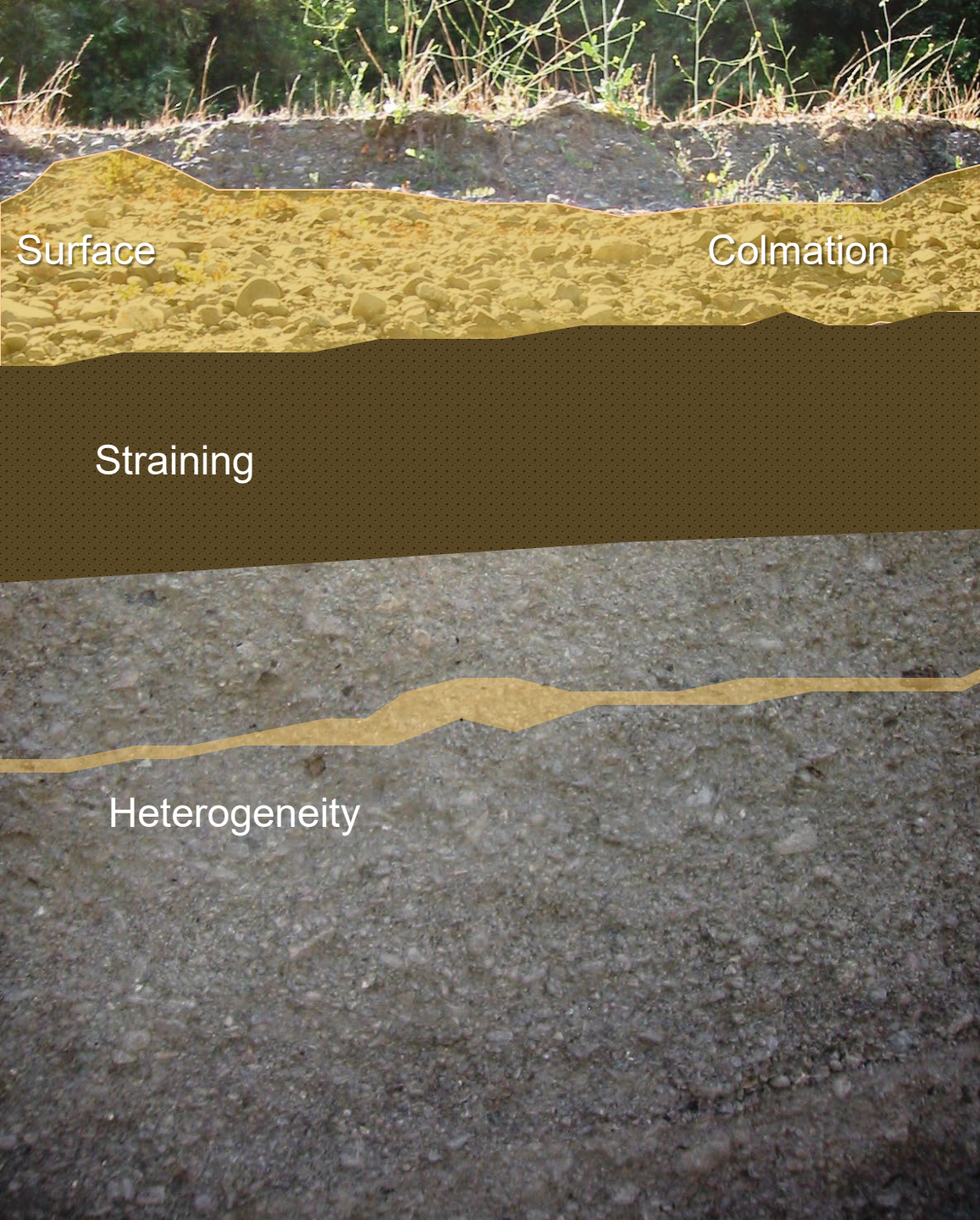
Crabby Creek, PA



Rio Calaveras, NM



Lake Tahoe, CA



Depth of hyporheic exchange is dependent on changes in K with depth

The vertical extent of hyporheic exchange also depends on the vertical distribution of K in the bed sediments. If low- K sediments are present at or near the surface, the volume and vertical extent of hyporheic exchange will be reduced. If low- K sediments are present deeper in the streambed sediments, hyporheic exchange may not be affected. In many fluvial systems, the presence, absence, and distribution of low- K sediments can change over short time scales.

Colmation is the clogging of sediments with fine-grained particles. Colmation can be focused at the surface (surface colmation) or vertically distributed in the bed (straining). Remember that D_{10} grain-size is often the most important for controlling K .

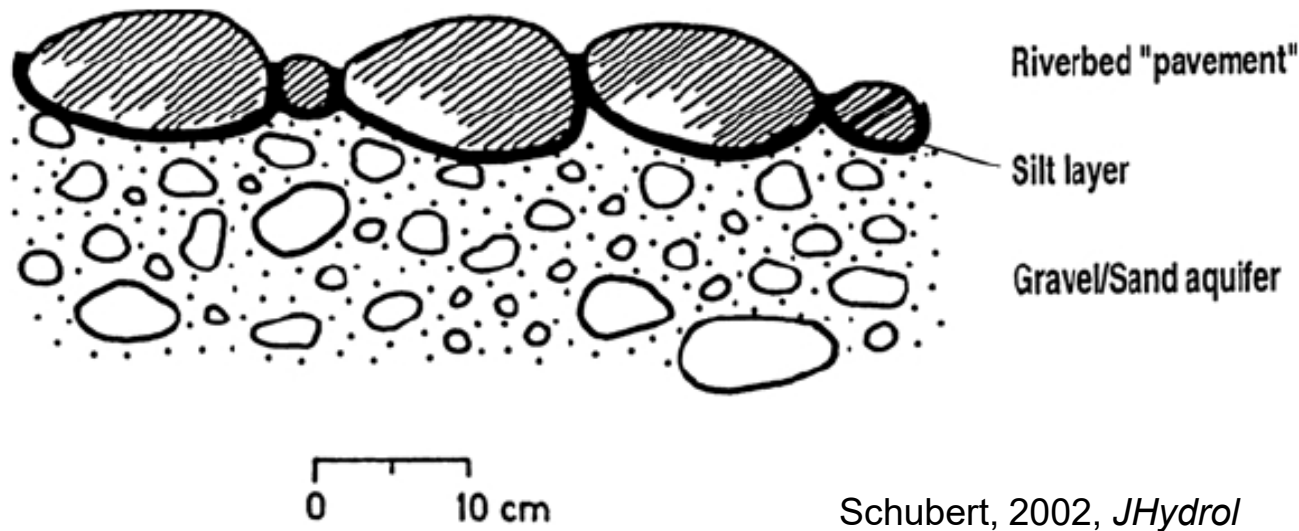
Examples of surface colmation



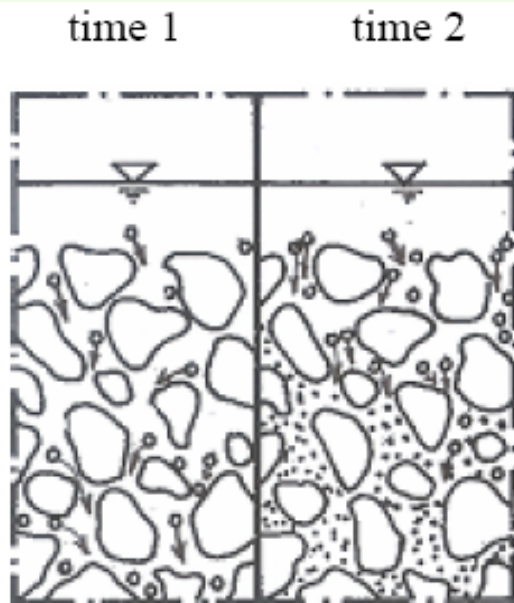
Here is an example of surface colmation. A layer of clayey silt deposited on a sand bed has dried after declining stream stage exposed the bed. Clay in the colmation layer has shrunk, causing the silt-clay mix to curl and expose the sand beneath. Here the colmation layer is perhaps 1 to 3 mm thick.



In this photo, a depression created by walking on the exposed but still wet streambed extends through the colmation layer to the sand and gravel beneath. Here the colmation layer appears to be about 1 cm thick.

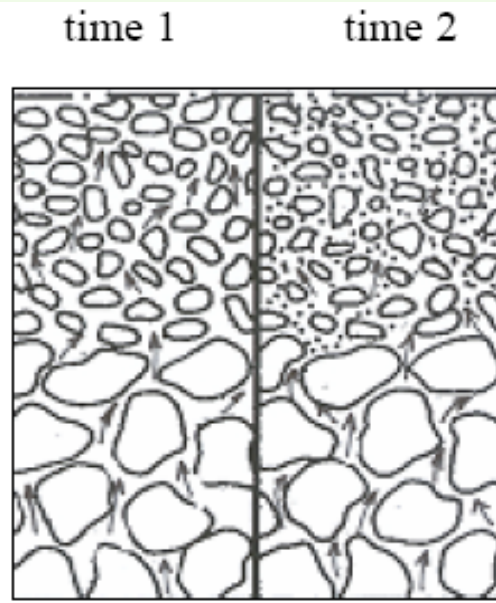


- **Colmation** can greatly limit hyporheic exchange **until the streambed is re-mobilized**
- Where surface armor is large and extensive and old, only a big flood can remobilize the bed
- Type of colmation depends on the ratio of the size of the clogging particles to the particles on the bed
- Biological processes can enhance colmation and make bed more cohesive (e.g., algal mat)
- Or they can reduce colmation and allow greater exchange



(B) **Internal** colmation

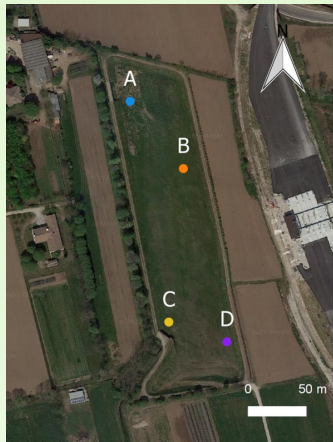
Straining



(C) **contact** colmation

Surface colmation

Velickovic, 2005, *Arch&Civil Eng.*



Infiltration basin in Italy

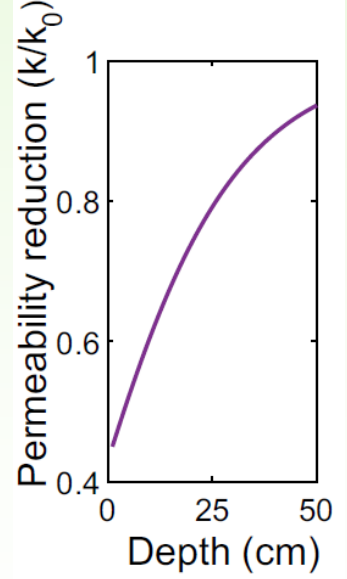
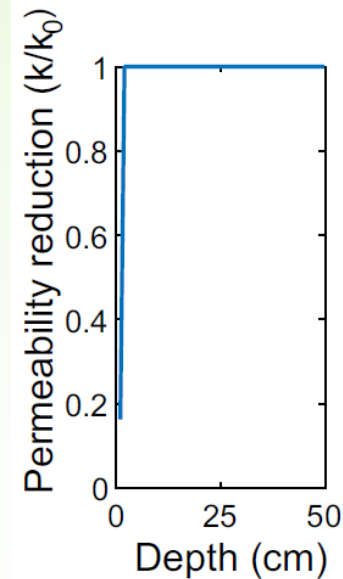
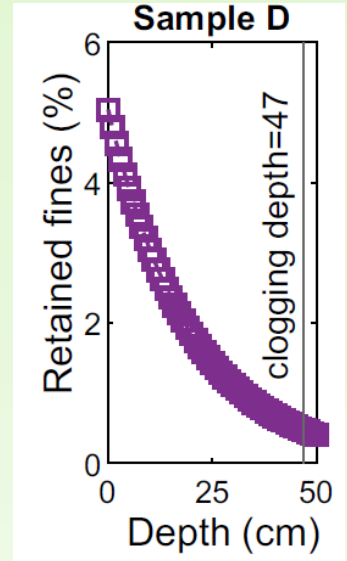
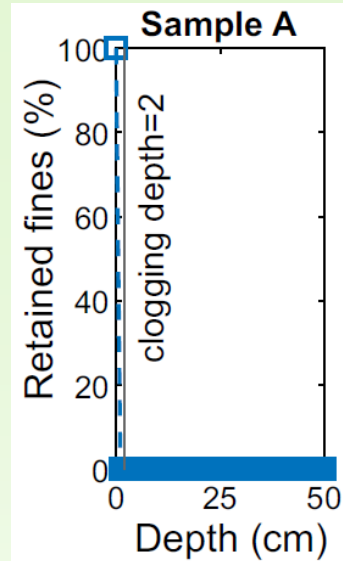
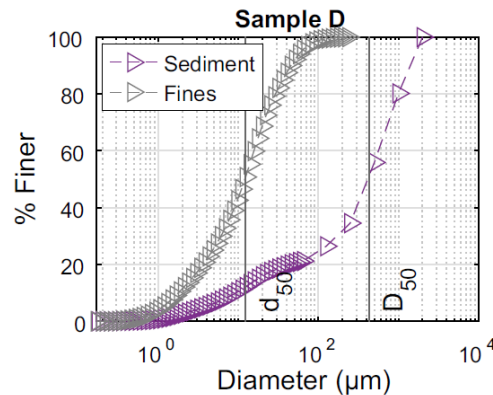
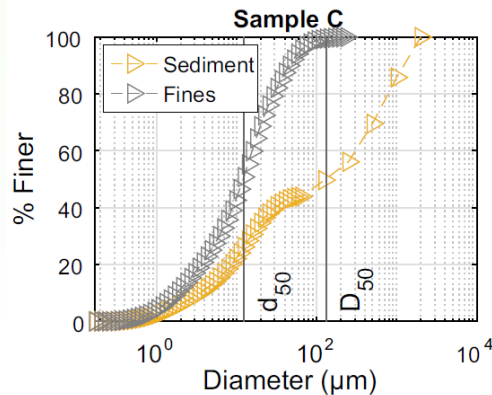
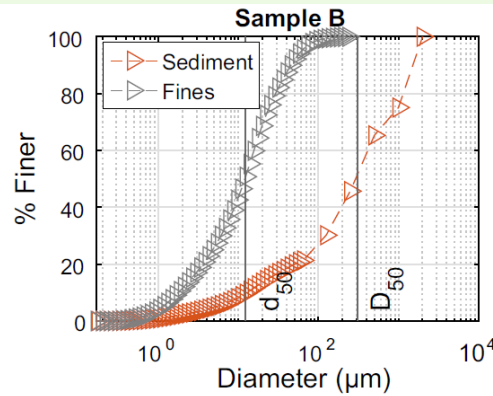
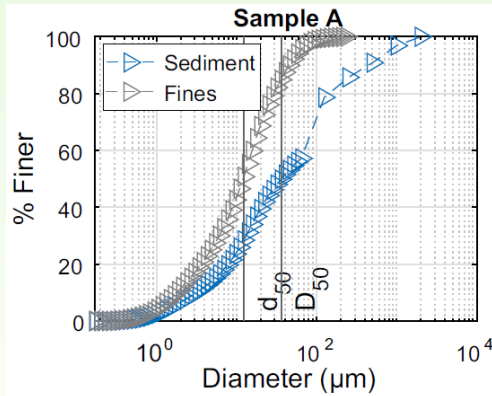
Modeled clogging depth and extent based on d_{50}/D_{50}

Infiltrating fines is clayey silt ($d_{50} = 12 \mu\text{m}$)

Sample A is medium silt ($D_{50} = 37 \mu\text{m}$)

Sample D is medium sand ($D_{50} = 0.4 \text{ mm}$)

Where fines infiltrate a silt bed, all the clogging is in the top 2 cm. Where fines infiltrate a sand bed, the fines are distributed across the top 50 cm of the bed. By 0.5 m depth, K is reduced by 92 percent.



Lippera et al., 2023, *HJ*

Why does colmation matter?

One reason is many drinking-water supplies induce flow from surface water to municipal wells (river-bank filtration or RBF). Colmation reduces the process.

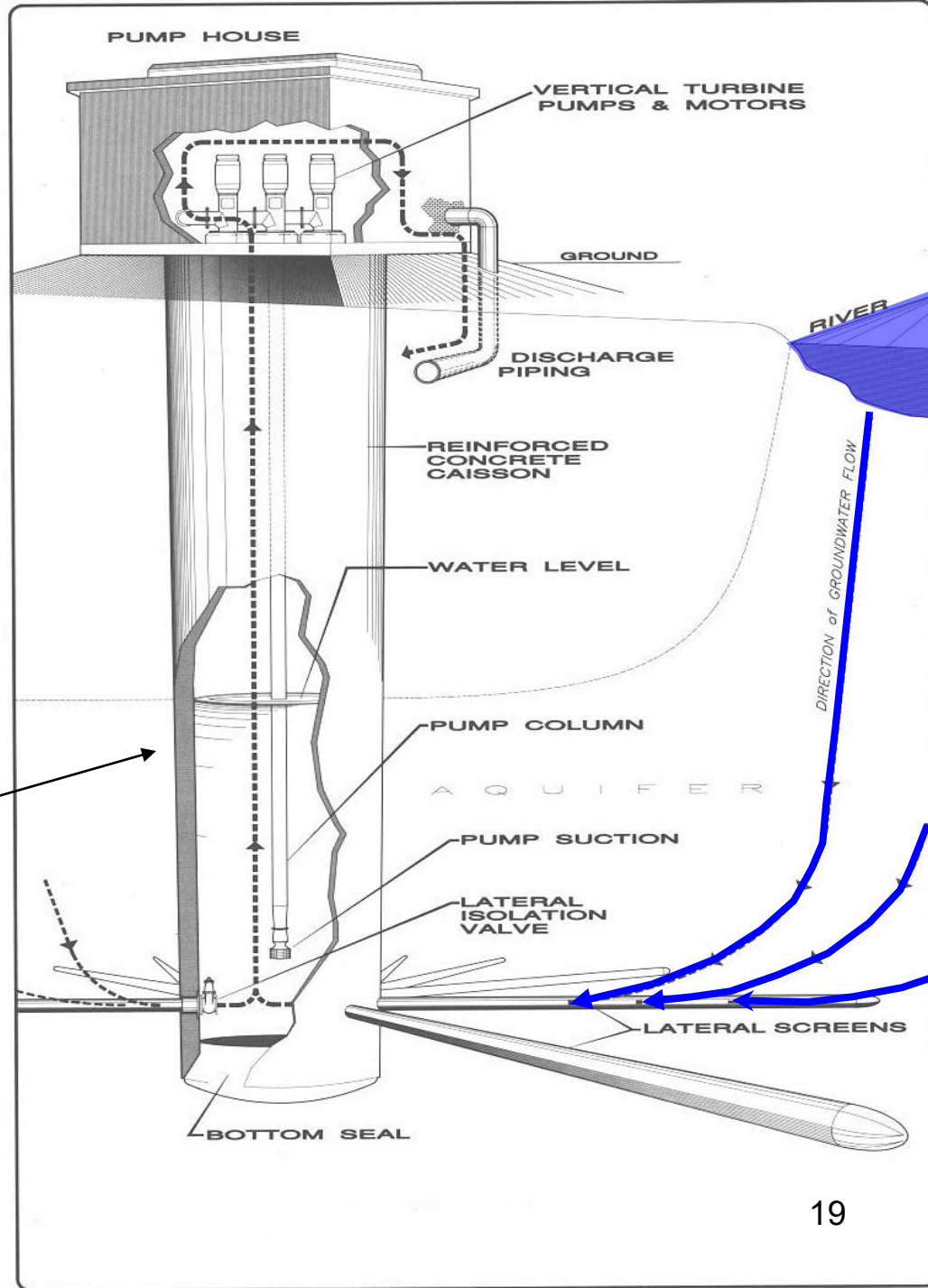
River bank filtration (RBF)

RBF supplies 16% of potable water for Germany

7% of potable water for The Netherlands

Ranney collector well

- Large diameter, high volume well
- Laterals extend beneath the river to induce flow from the river to the well
- Water-treatment requirements are often less restrictive because we assume bacteria and viruses in surface water are filtered by the sediments before they reach the well.

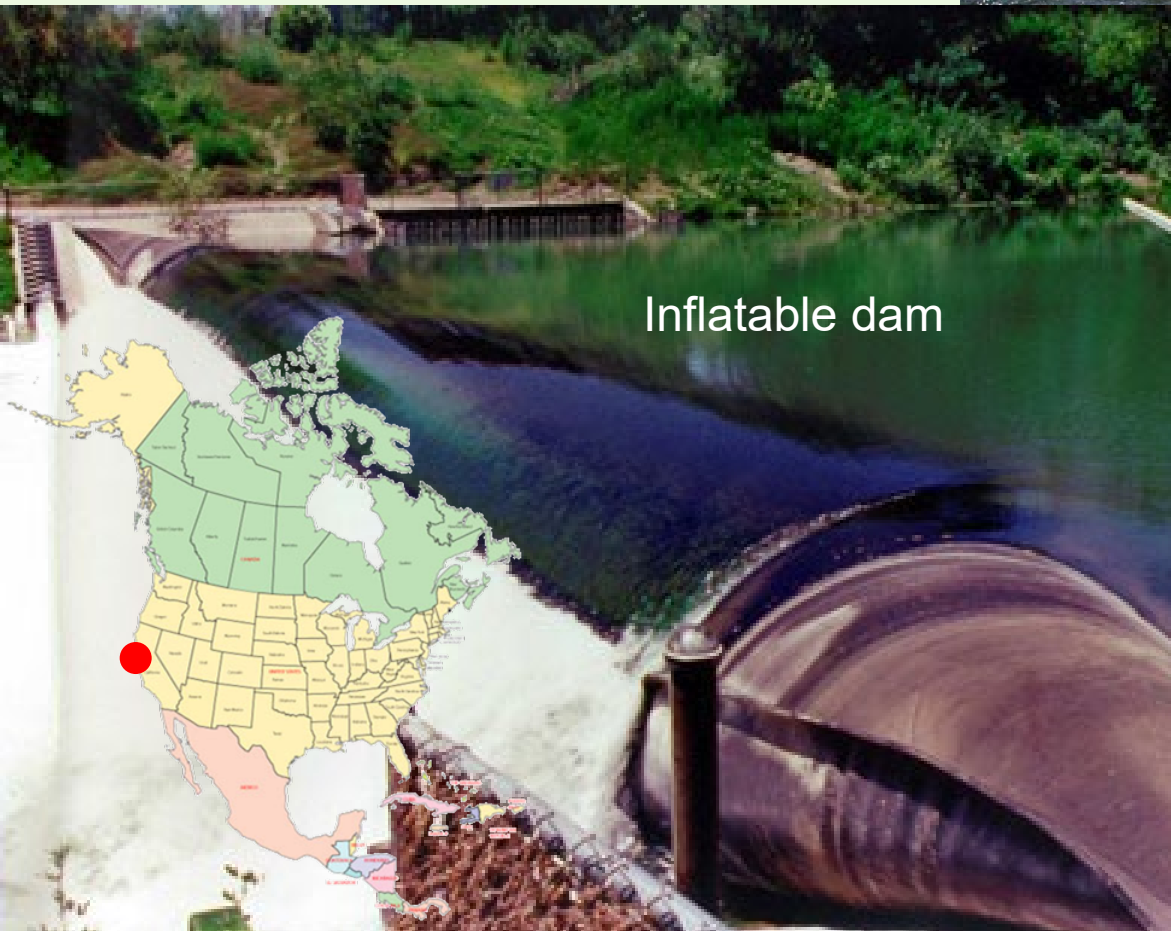


Sonoma County Water Agency, Russian River, CA

- Produce 300,000 m³/day
- Inflatable dam raises river stage 3 m
- Still can't get enough water because the riverbed is becoming clogged



Russian River



Inflatable dam



4-m diameter
Ranney collector well



Induced infiltration depends on algae accumulation, basin stage, and accumulation of fine-grained sediments

(Here we are using seepage meters to determine areal distribution of seepage)

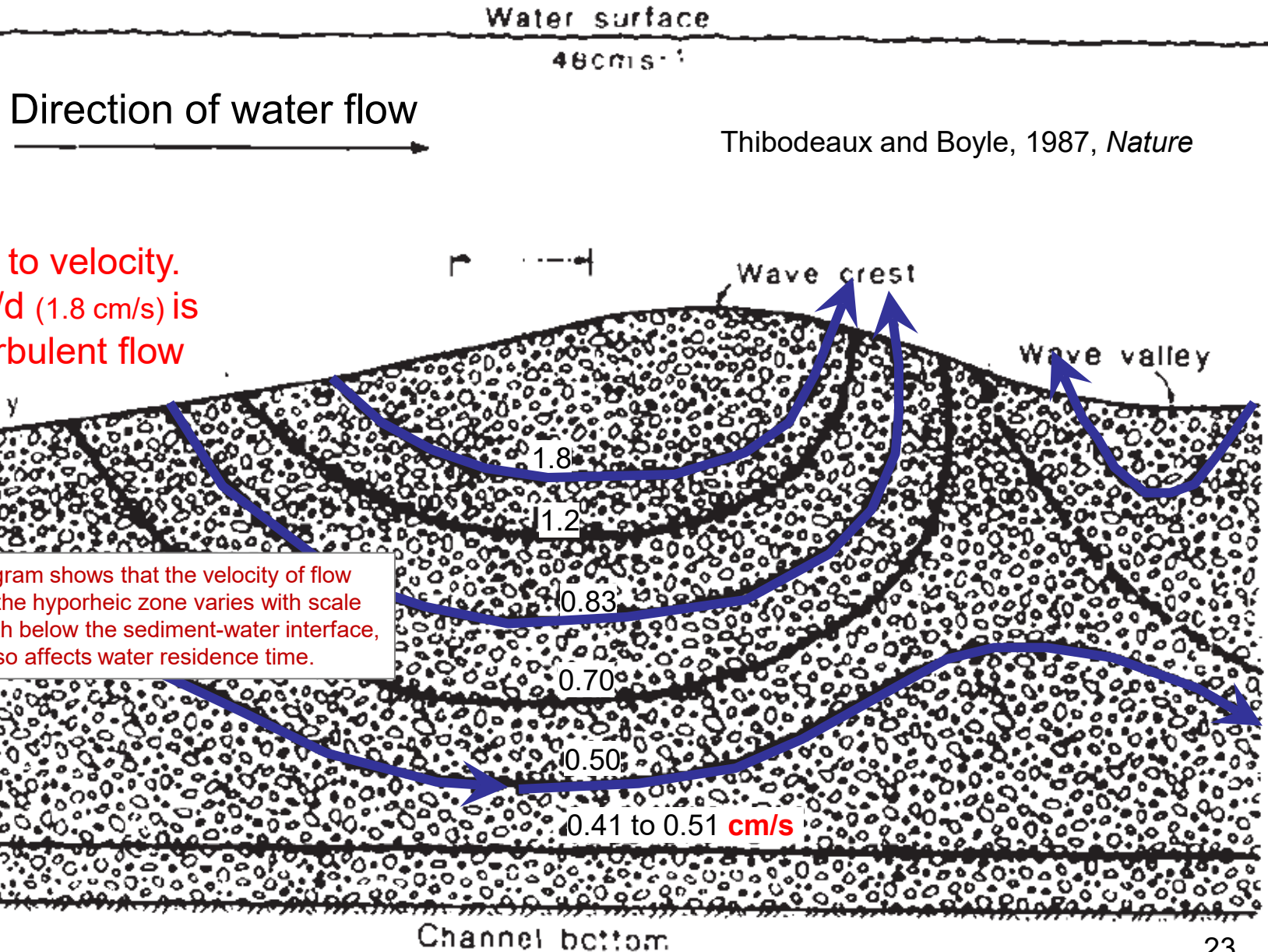


Dealing with colmation

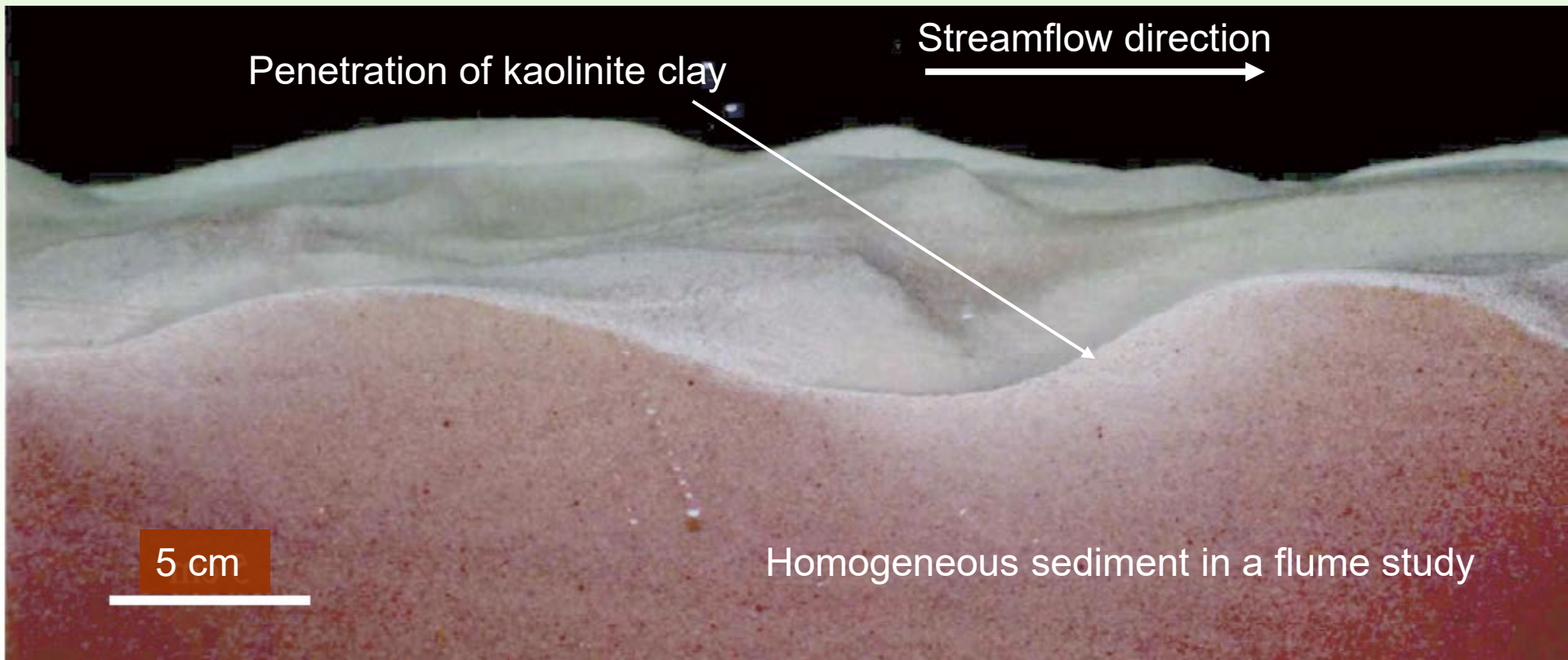


This infiltration basin was drained because infiltration rates had decreased. The scraper in the photograph is removing the fine-grained sediments that had accumulated at the basin surface. Following removal, the basin will be filled again with river water and infiltration to ground water and the nearby pumping well will be much faster until the bed once again becomes clogged..

Bedform scale – Bedforms drive convective exchange in bed sediments



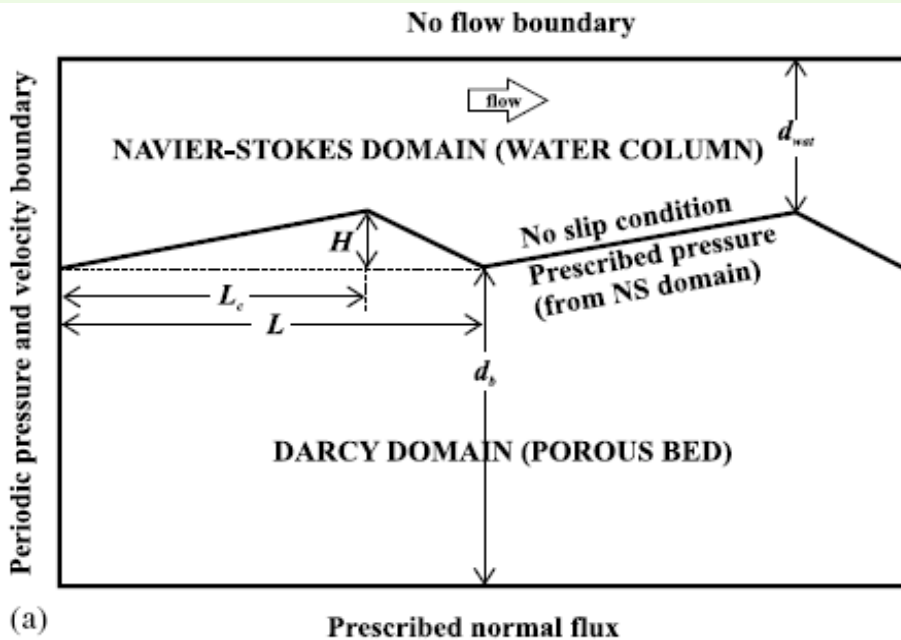
Sediment and chemicals are transported along hyporheic flowpaths along with the water



Note that flow here is from left to right. White clay particles penetrate 5 cm or more into medium-grain sand, shown here as pink. This indicates that sediment, and any chemicals sorbed to sediment particles, can move through hyporheic sediments along with the water.

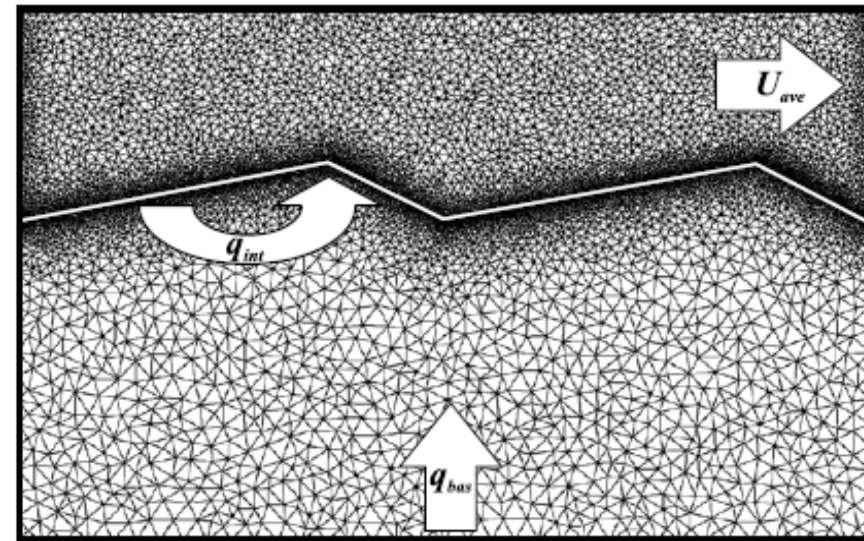
Modeling of hyporheic exchange with changing GW discharge

- Modeled laminar flow in porous medium and turbulent flow in stream
- Looked at the influence of bedform on inducing hyporheic exchange
- Held surface-water velocity constant



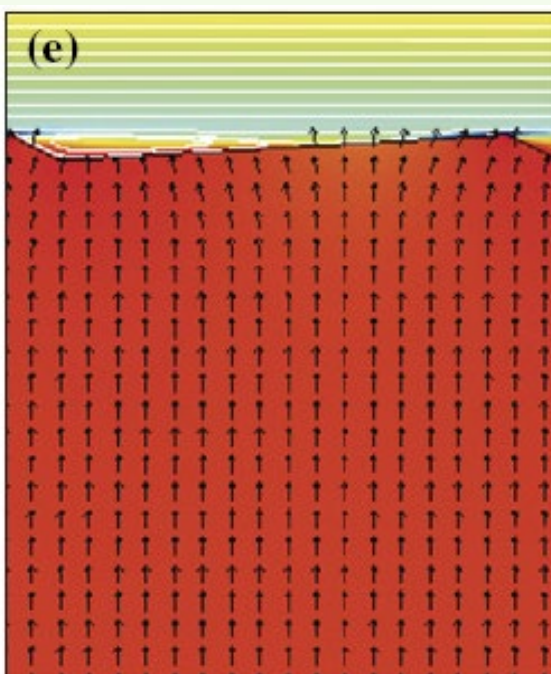
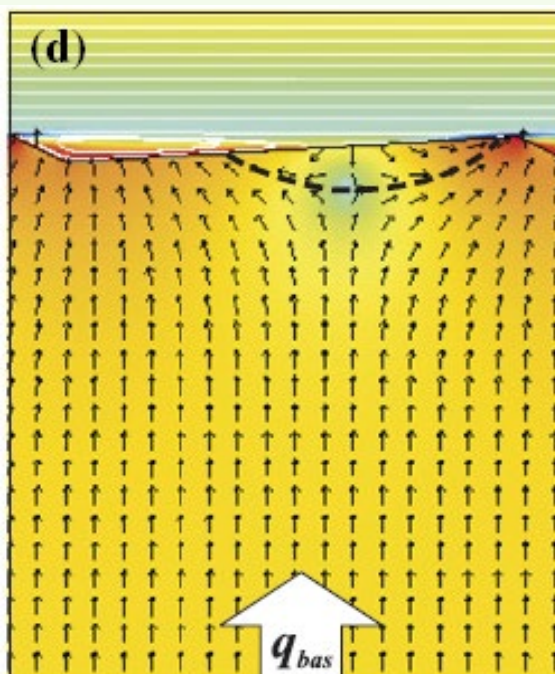
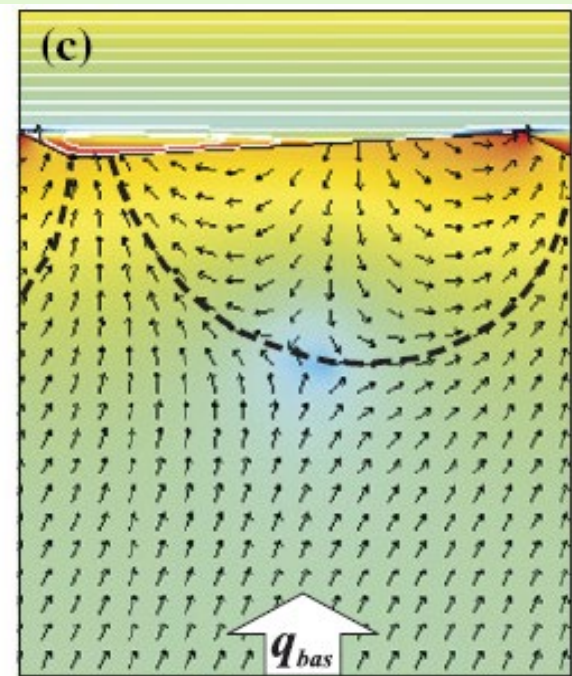
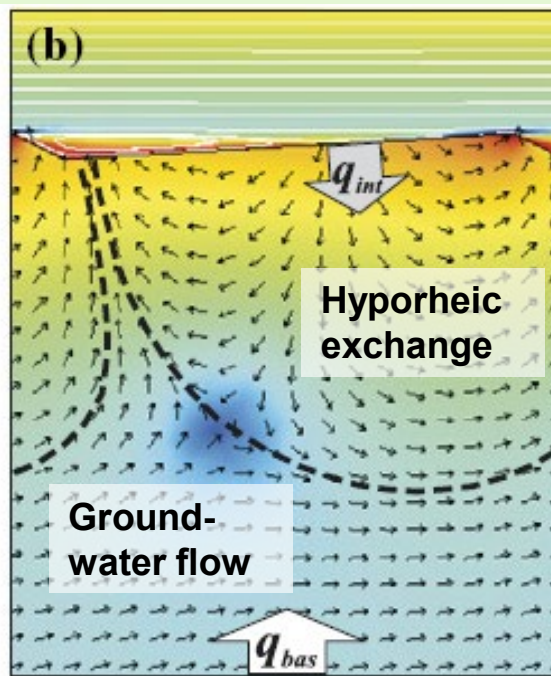
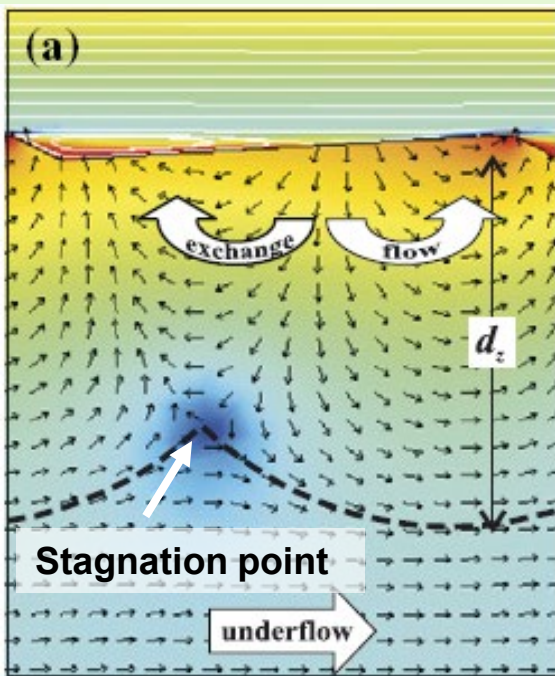
Periodic pressure (plus pressure drop dp) and velocity boundary

(b)



Bayani Cardenas has contributed much to the hyporheic literature. In this paper, he and his thesis advisor, John Wilson, coupled a surface-water hydrodynamics model with a groundwater-flow model to study the effect of groundwater discharge on the boundary between hyporheic water and unaltered groundwater.

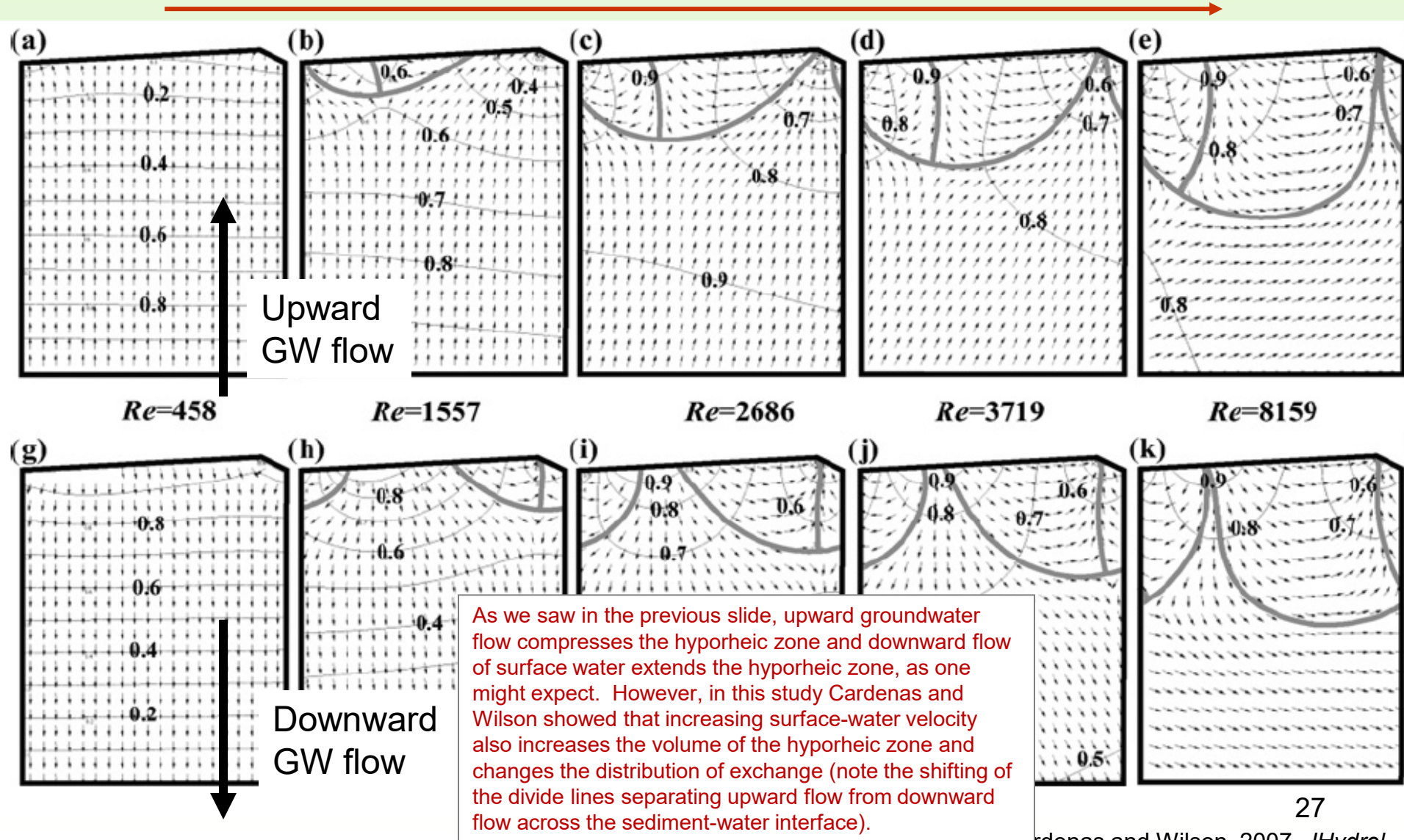
Cardenas and Wilson, 2006,
JHydrol



- Depth of hyporheic zone decreased with increasing GW discharge
- Line dividing hyporheic exchange from GW flow represents a steep gradient for physical, chemical, and biological parameters (a hyporheic ecotone of sorts)

Hyporheic zone also depends on surface-water velocity, K , and direction of seepage

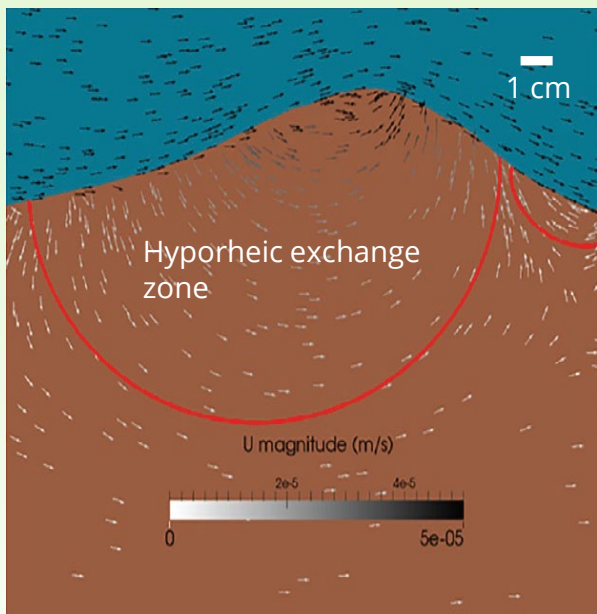
Increasing velocity



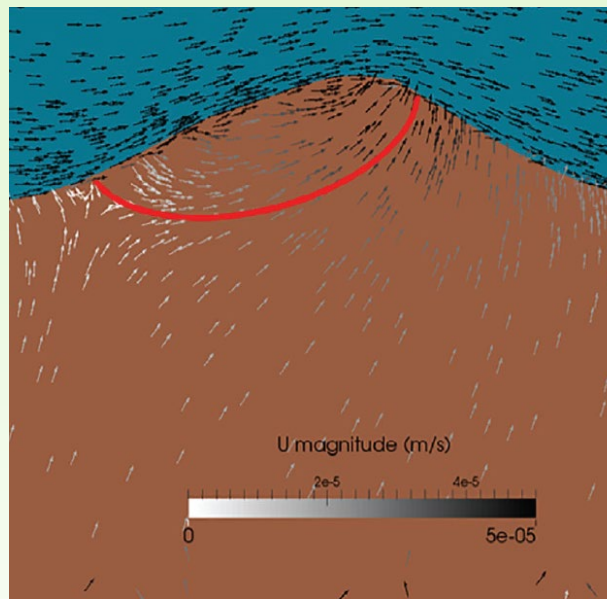
Hyporheic zone also depends on where the water that enters the porous media ends up

Authors used OpenFOAM to simulate both surface-water and groundwater flow domains at a bed ripple scale.

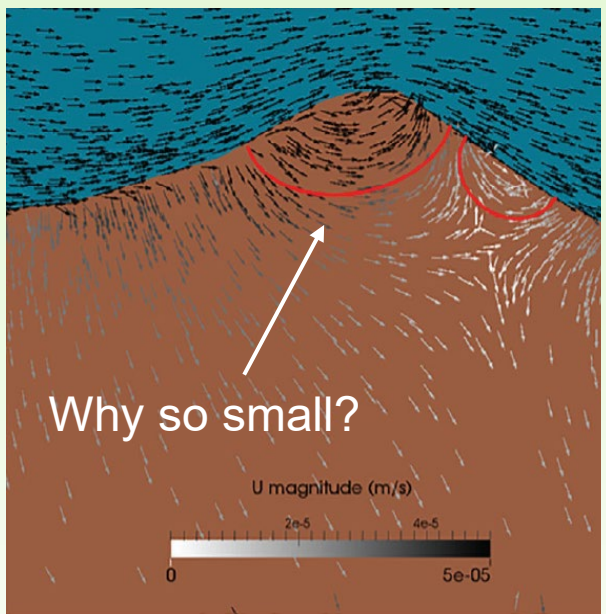
No vertical gradient



Upward flux at +49 cm/day

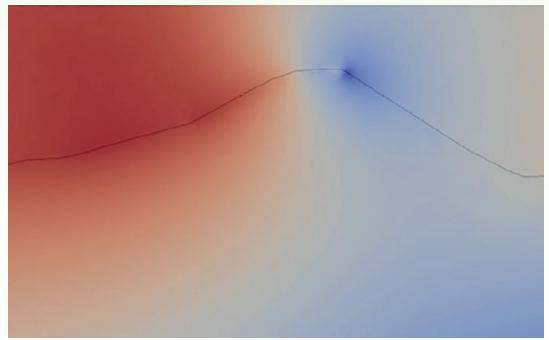
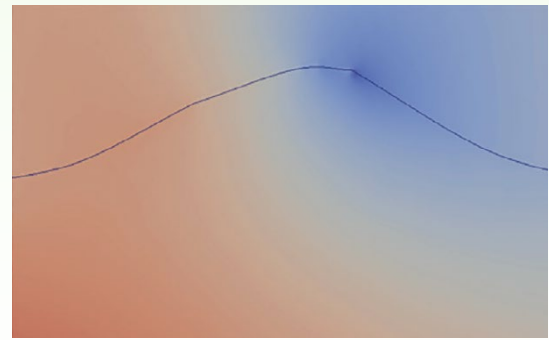


Downward flux at -49 cm/day

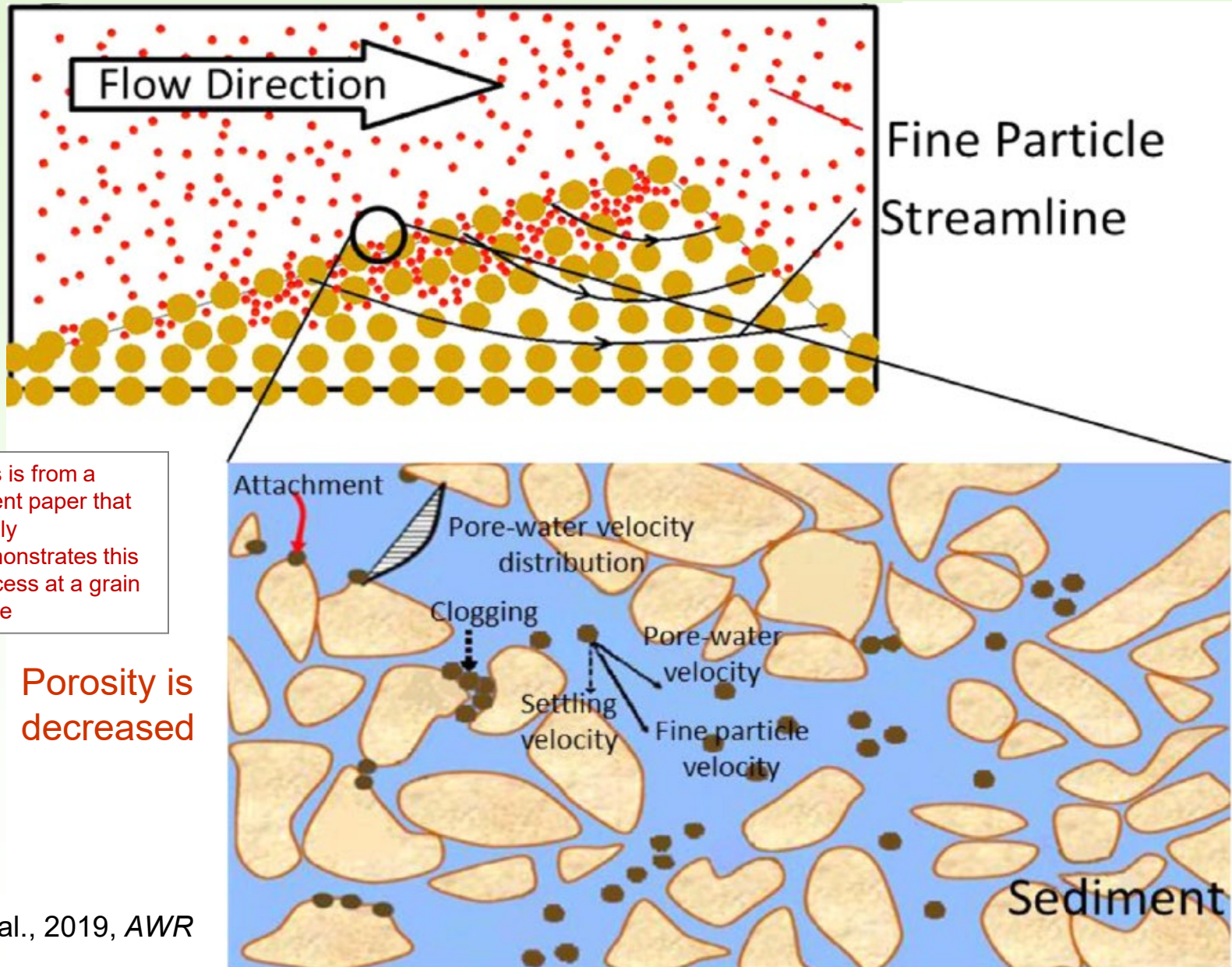


More powerful, more recent modeling tools lead to better understanding

The volume of the hyporheic zone is reduced (compressed) by upward flux through the groundwater flow domain, which makes logical sense. But why is the hyporheic zone for downward flux smaller than the neutral condition? It's because more of the groundwater flow paths continue downward through the groundwater flow domain rather than flowing back into the stream.



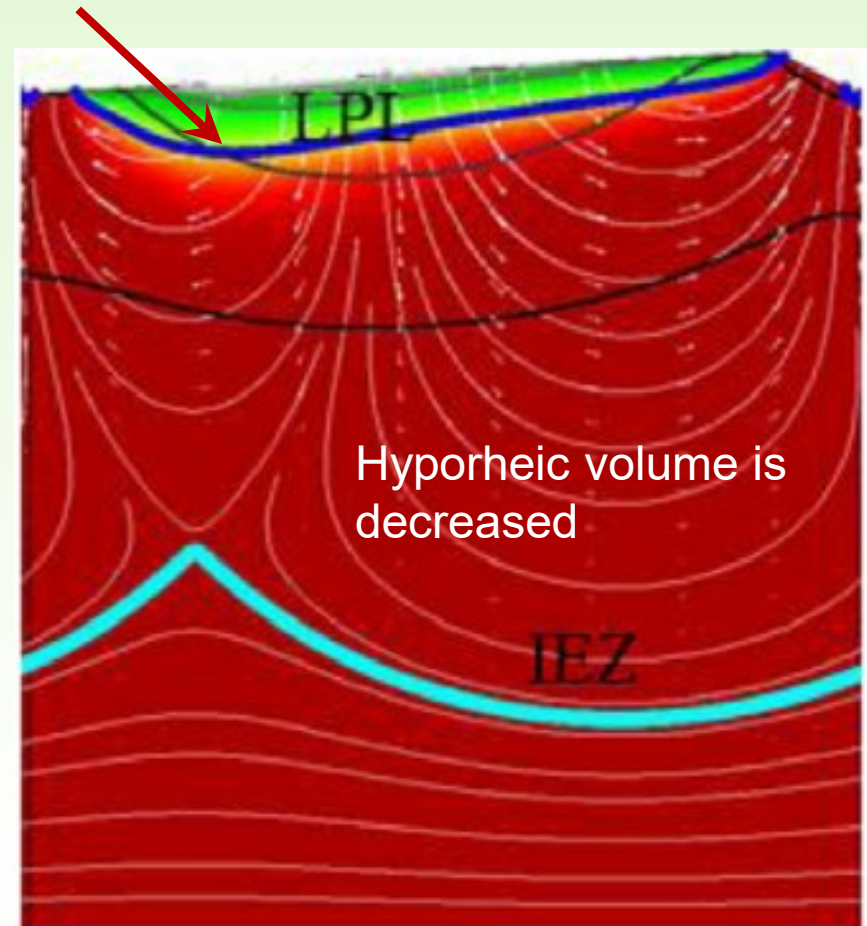
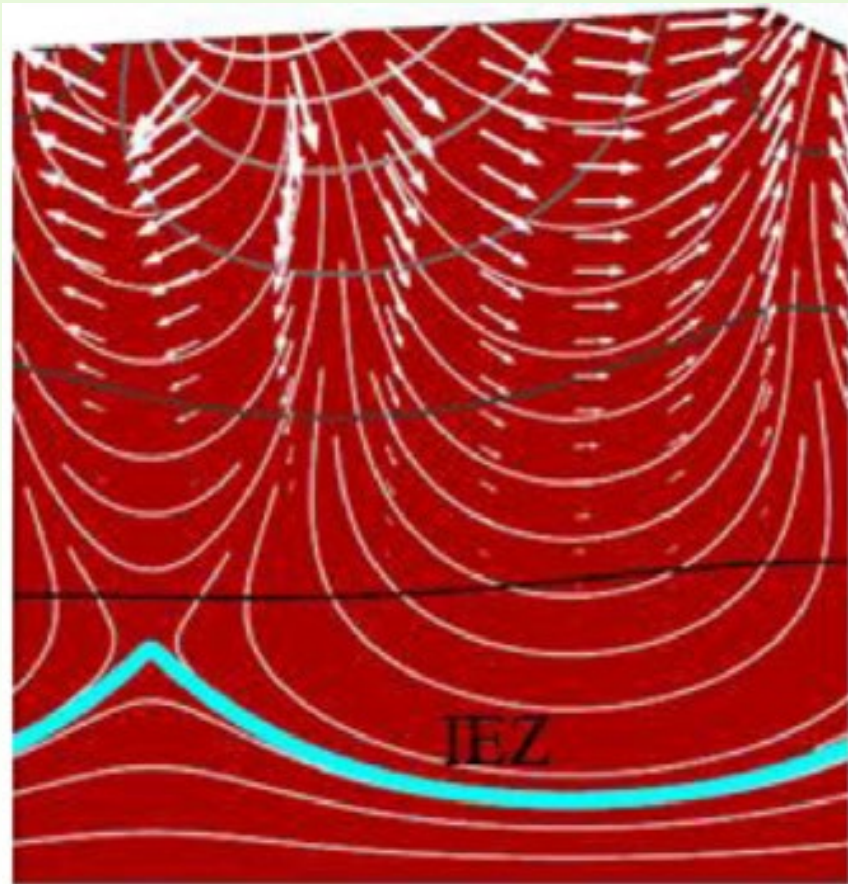
But, as mentioned before, clogging (colmation) can reduce hyporheic exchange



Near total clogging could occur in 1 to 20 hours

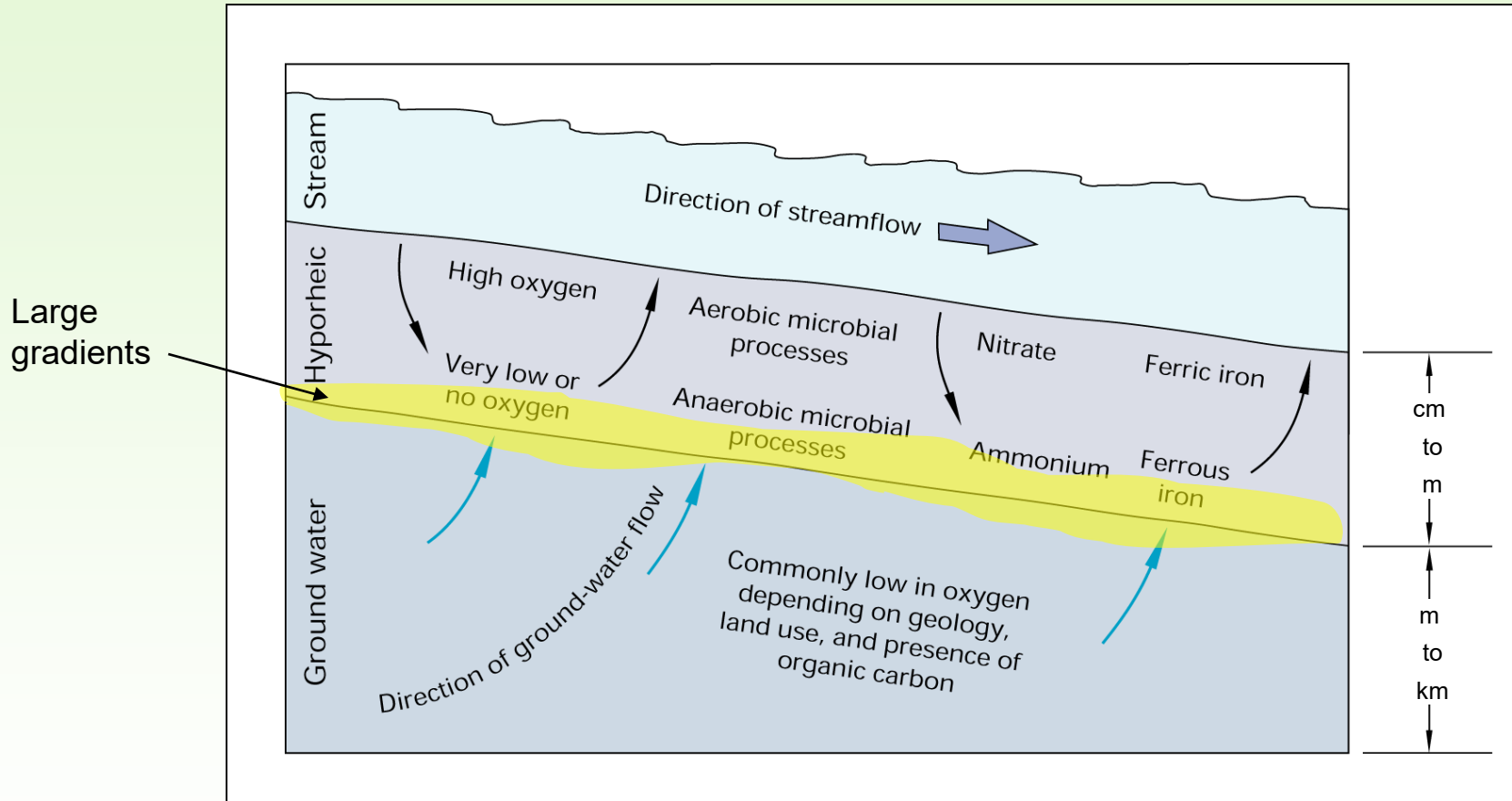
Clogging yes, but how long does it take? These results indicate clogging after a flushing flow that resets a sediment bed could occur within a day.

Low permeability layer



Jin et al., 2019, AWR

Biological and geochemical (ecological) perspective



Microbial activity and chemical transformations commonly are enhanced in the hyporheic zone compared to those that take place in ground water and surface water. This diagram illustrates some of the processes and chemical transformations that may take place in the hyporheic zone. Actual chemical interactions depend on numerous factors including aquifer mineralogy, shape of the aquifer, types of organic matter in surface water and ground water, and nearby land use.

Spawning redds

Another impetus that is driving much research by ground-water hydrologists, fluvial geomorphologists, and ecologists is whether fish create their spawning redds in areas of ground-water discharge. Whether they do or not seems to be species dependent. These measurements are particularly difficult to make because the streambed substrate often is very coarse. This is a rapidly growing research area.



- Alexander and Caissie. 2003. *Ground Water*
Brown and Ford. 2002. *River Research and Applications*
Moir et al. 2002. *Geomorphology*
Morrison et al. 2002. *Journal of Hydrology*
Soulsby et al. 2001. *Regulated Rivers: Research & Management*
Baxter and Hauer. 2000. *Journal of Fisheries and Aquatic Science*
Baxter and McPhail. 1999. *Canadian Journal of Zoology*
Garrett et al. 1998. *Journal of Fisheries Management*
Pitlick and Van Steeter. 1998. *Water Resources Research*
Ridgway and Blanchfield. 1998. *Ecology of Freshwater Fish*

Field example of reach-scale hyporheic exchange

Rosenberry & Pitlick, 2009, *HP*

This study demonstrates the complexity of GW-SW exchange in coarse-grained fluvial settings.

Diversion dam

D

C

B

A

Direction of flow



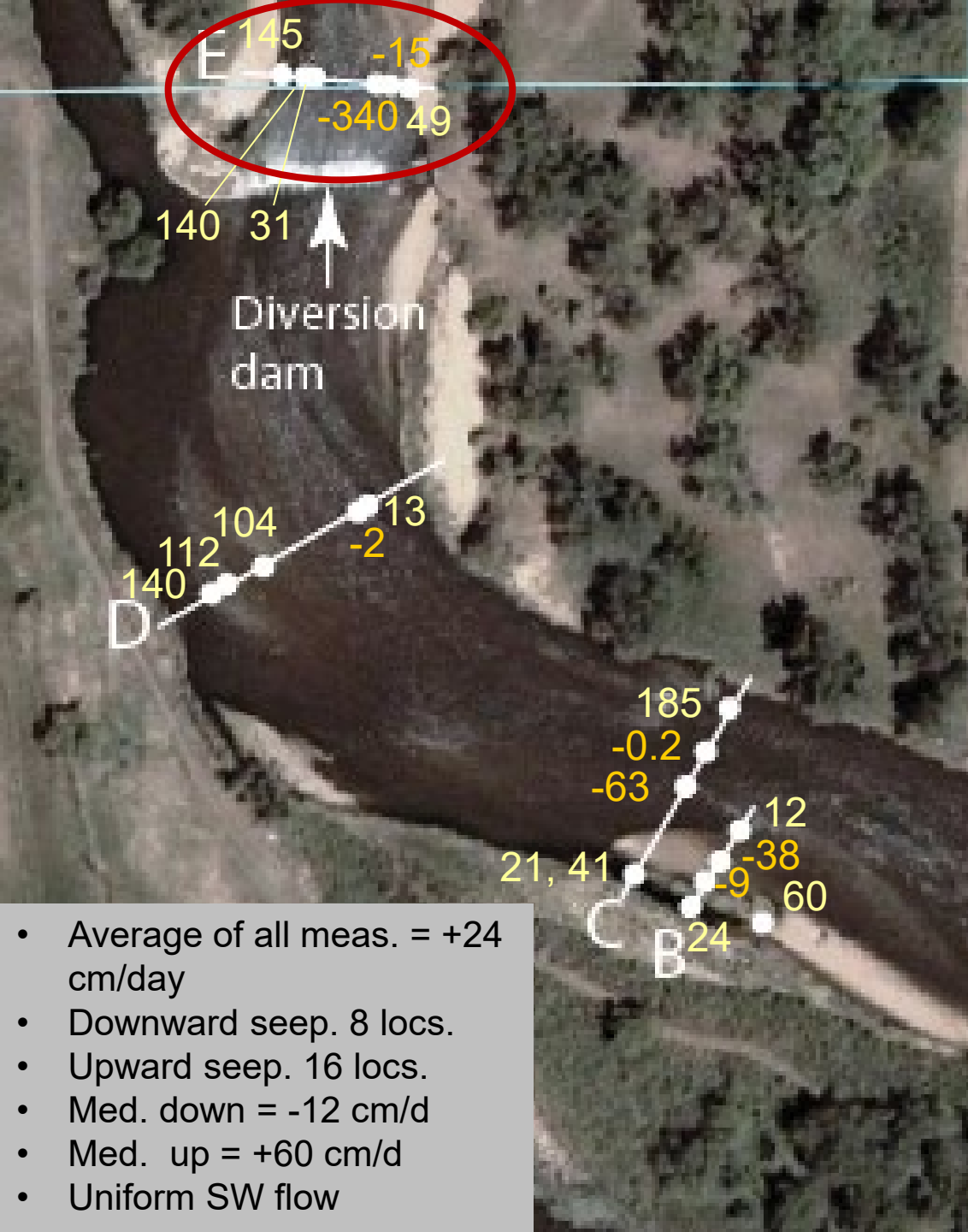
Image © 2008 DigitalGlobe
USGS

First, we had to develop a seepage meter that could be used in flowing water. The cylinder is streamlined to reduce drag and a bag shelter removes hydraulic effects from the seepage bag. We will talk more about these modifications in another lecture.

Low-profile seepage cylinder

Bag shelter





- Seepage is fast
- Spatial variability is large

Values are vertical flow across the riverbed expressed in cm/day. Positive values indicate upward flow and negative values indicate downward flow.

- Average of all meas. = +24 cm/day
- Downward seep. 8 locs.
- Upward seep. 16 locs.
- Med. down = -12 cm/d
- Med. up = +60 cm/d
- Uniform SW flow

Bed topography controls seepage

-340

-15

68
cm/s

u

The top meter is on a submerged bar. The meter where seepage is -15 is in the thalweg. Surface-water current is 68 cm/s (if flow is much faster then the meters get flushed out of the sediment. Faster flow at the bar is due in part to the bar extending into the surface-water flow field and water being pushed into the bar sediments.

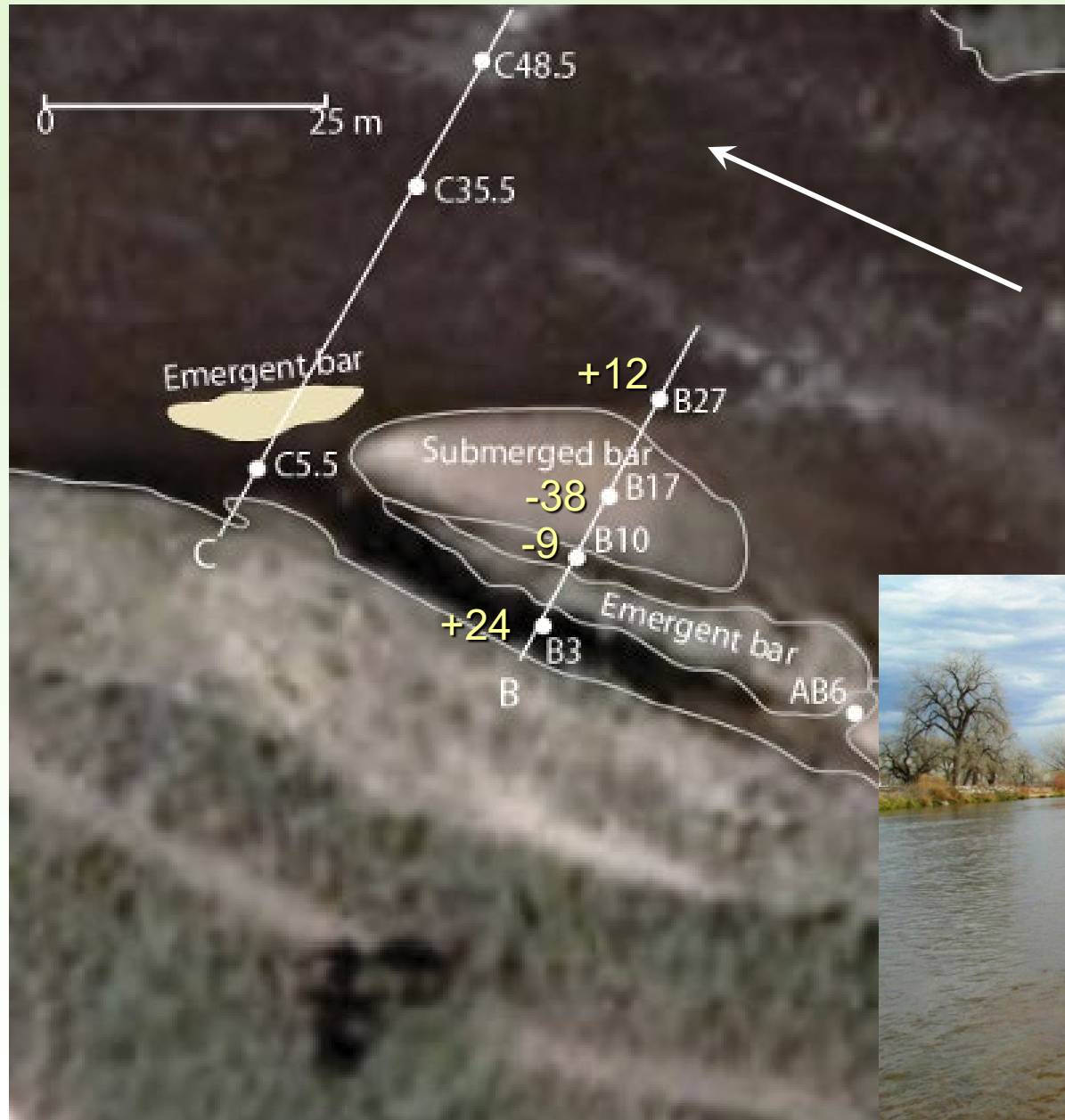


The same meter

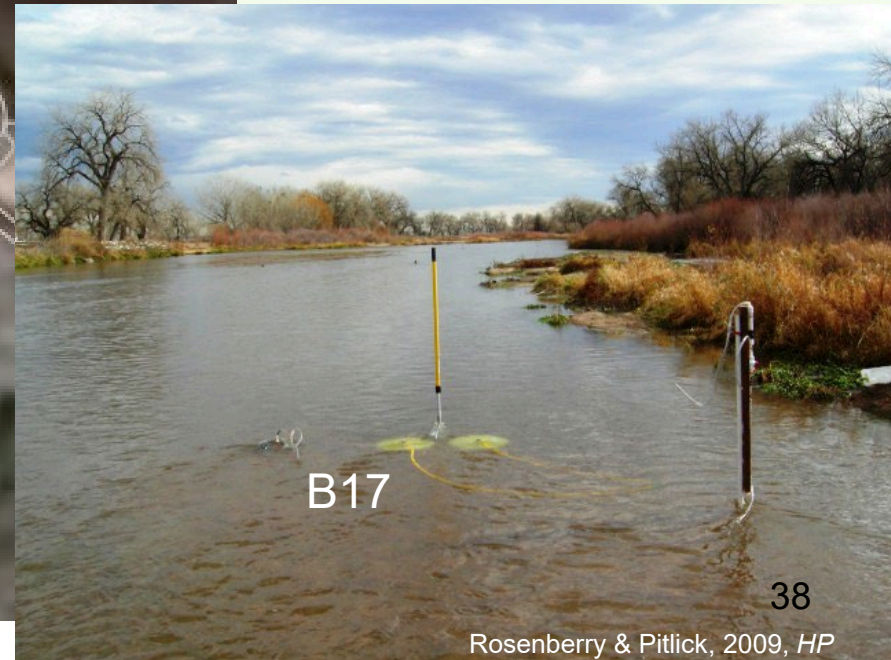


But downward seepage (at both locations) also is due to larger-scale exchange caused by two braids of the river being at different elevations. Water is seeping downward through the bed in the channel to the left and flowing toward the channel just visible at the right edge of the photograph.

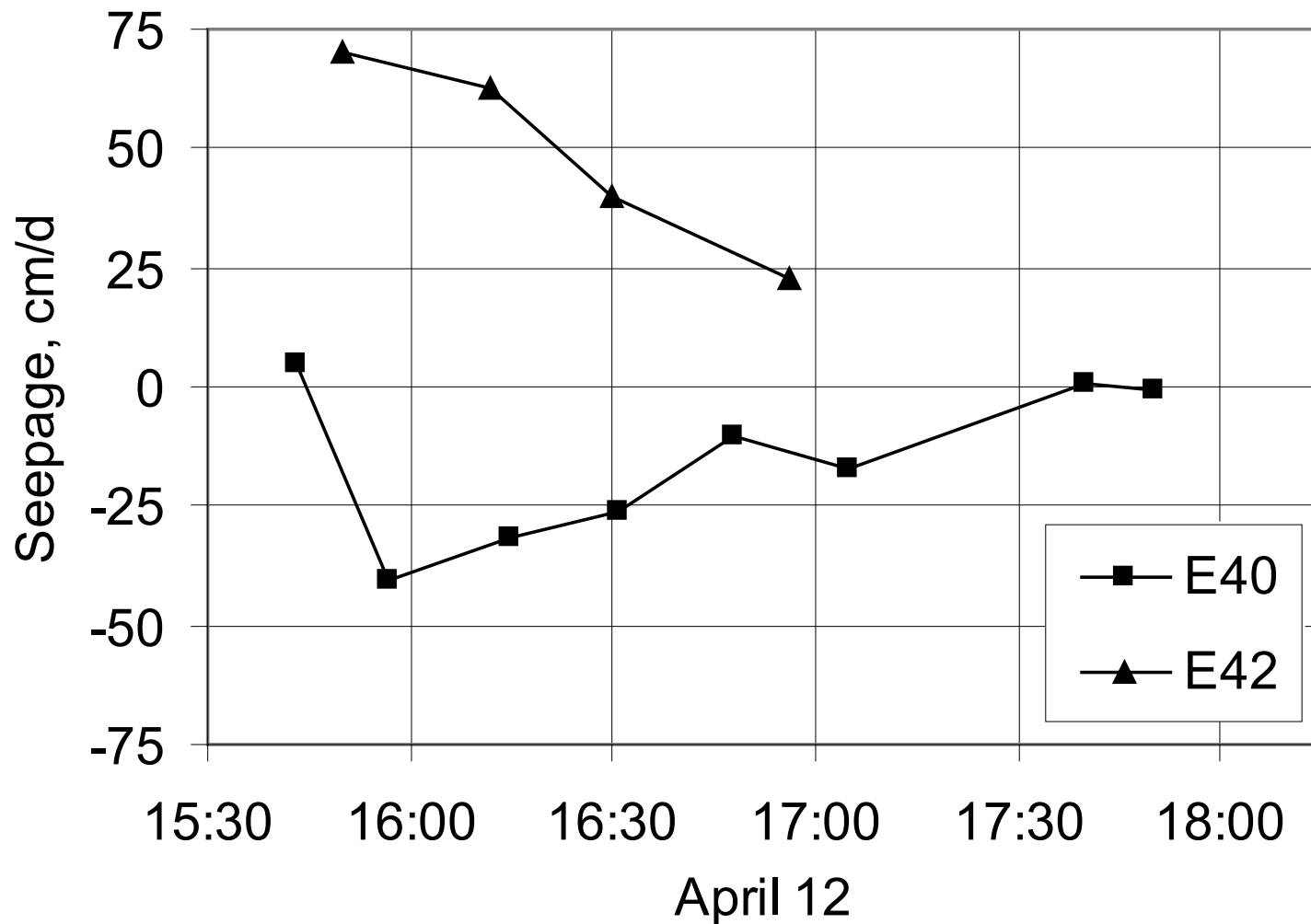
Bed topography controls seepage



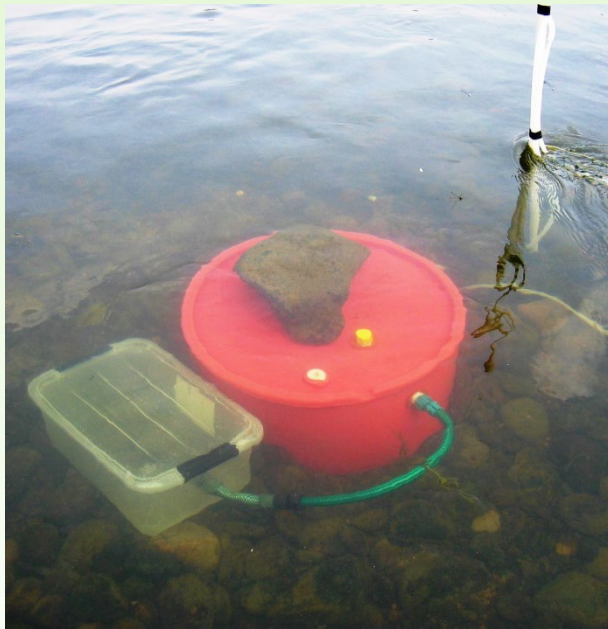
Here is another example of bed topography controlling distribution of seepage. Seepage is downward where water is shoaling over a submerged bar, but seepage is upward in slightly deeper water at location B27 (and also at B3 and C5.5) The inset photo shows measurements of seepage (two yellow cylinders), sediment transport (meter placed on the bed just upstream of the seepage meters), hydraulic gradient (manometer near the right side of the photo), and use of a piezoseep that will be discussed later.



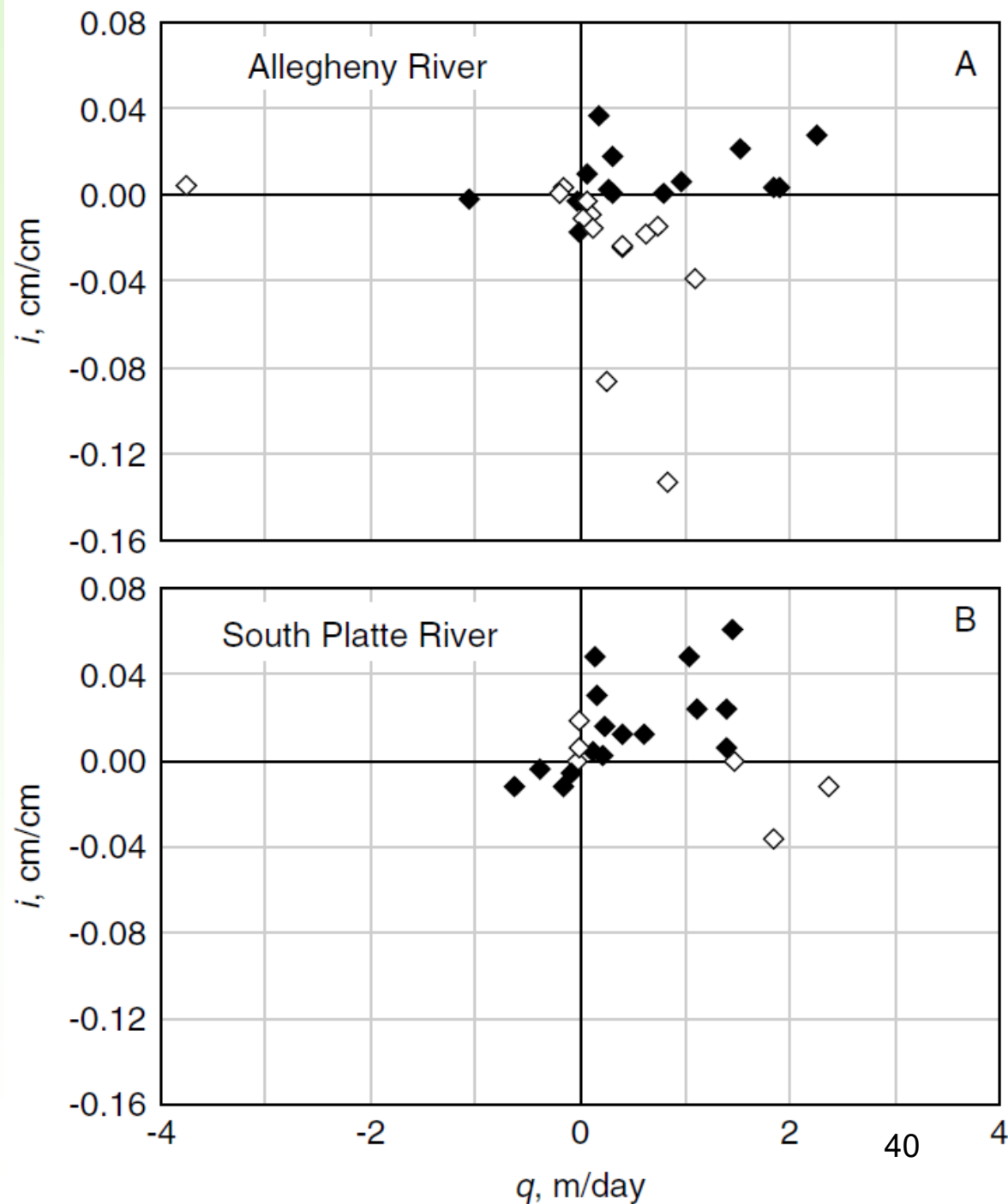
And exchange can vary substantially over time



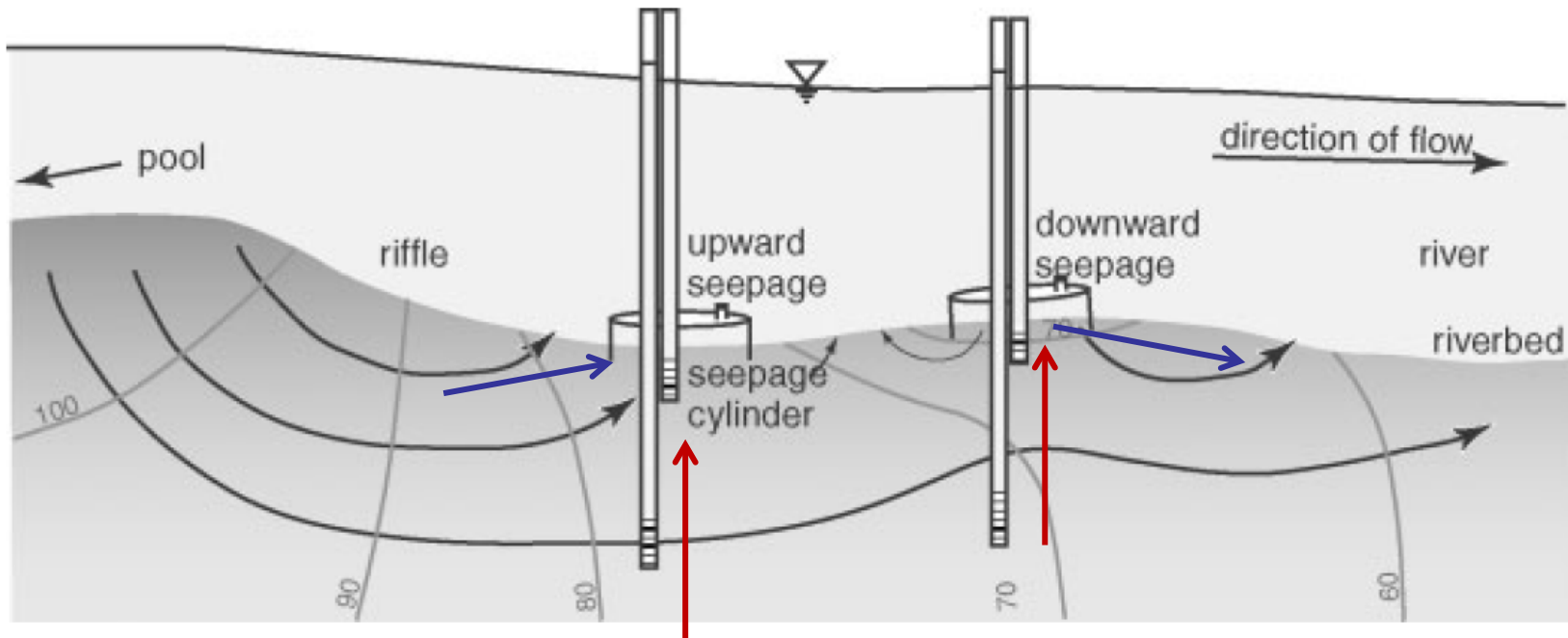
q and i might not
correlate very well in
hyporheic settings



The open diamonds are locations where the seepage meter indicated flow in one direction and the adjacent piezometer indicated the potential for flow in the opposite direction. Why?



Riverbed heterogeneity may result in different instruments measuring processes at different scales



Seepage is primarily horizontal

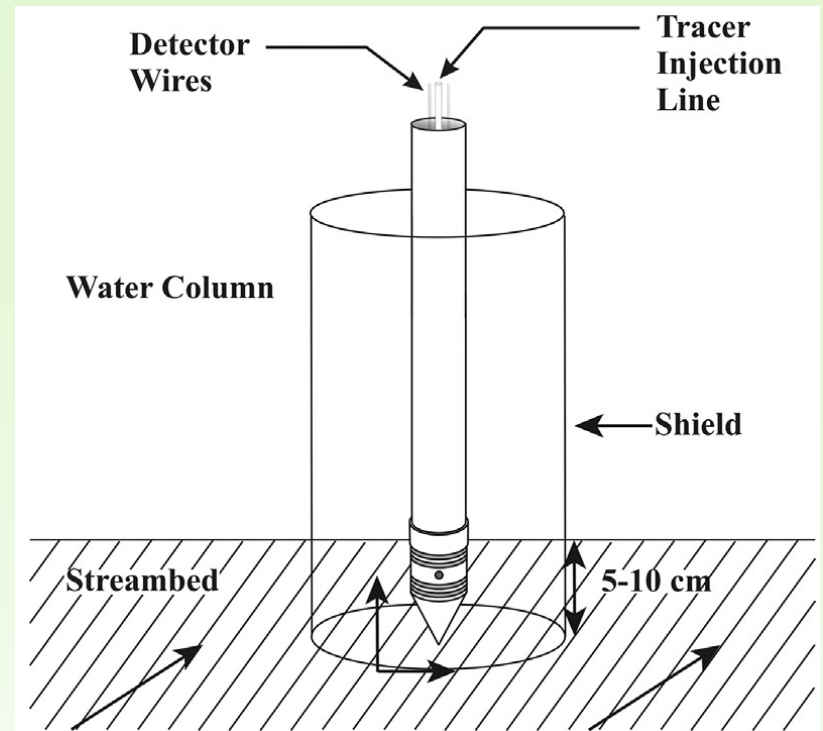
Gradient (i) measured along a vertical axis

When we measure a hydraulic gradient with a piezometer installed in a streambed, the gradient is measured along a vertical axis (unless the piezometer is installed at an angle). However, seepage measured with a seepage meter is often in response to flowpaths that are not vertical. In some settings, interpretations of seepage based on these two methods can be quite different, including settings where hydraulic gradients and seepage meters can indicate flows in opposite directions. Such is the case in the seepage meter and piezometer shown on the right.

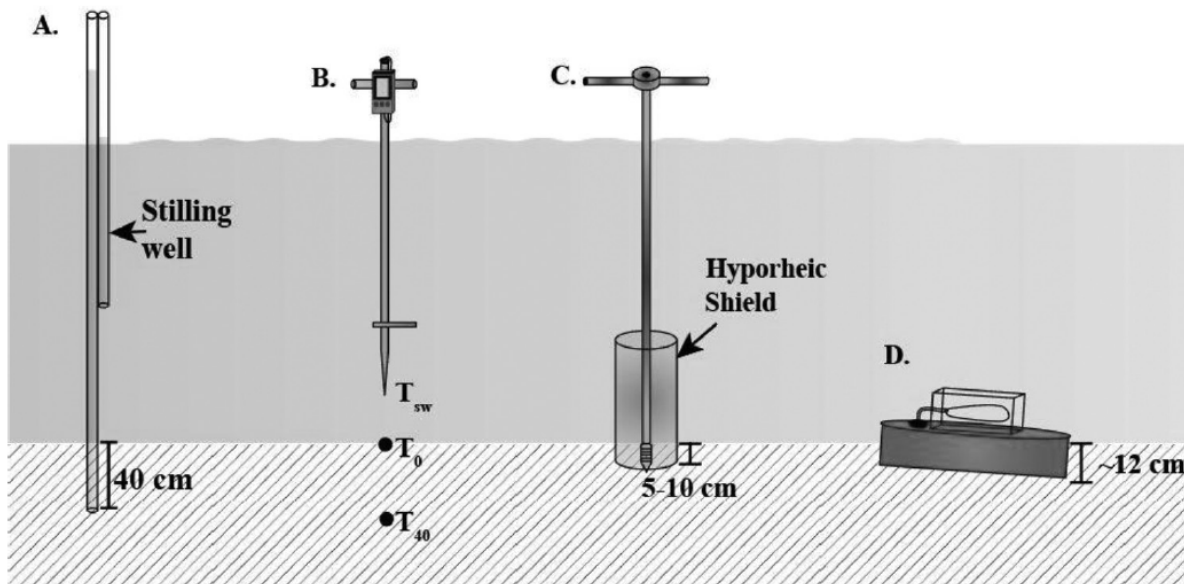
Streambed point-velocity probe

Modified to measure only vertical component of flux

They deployed a point-velocity probe that measures velocity based on timed tracer injections, a device typically used to measure flow either in boreholes or screened intervals of wells. They placed a metal cylinder, open on both ends, to block horizontal flow and allow measurement of only the vertical component of flow. They went on to compare this tool with other methods presented in a paper published in 2020. Those results are presented in a later lecture on methods for quantifying GW-SW exchange.

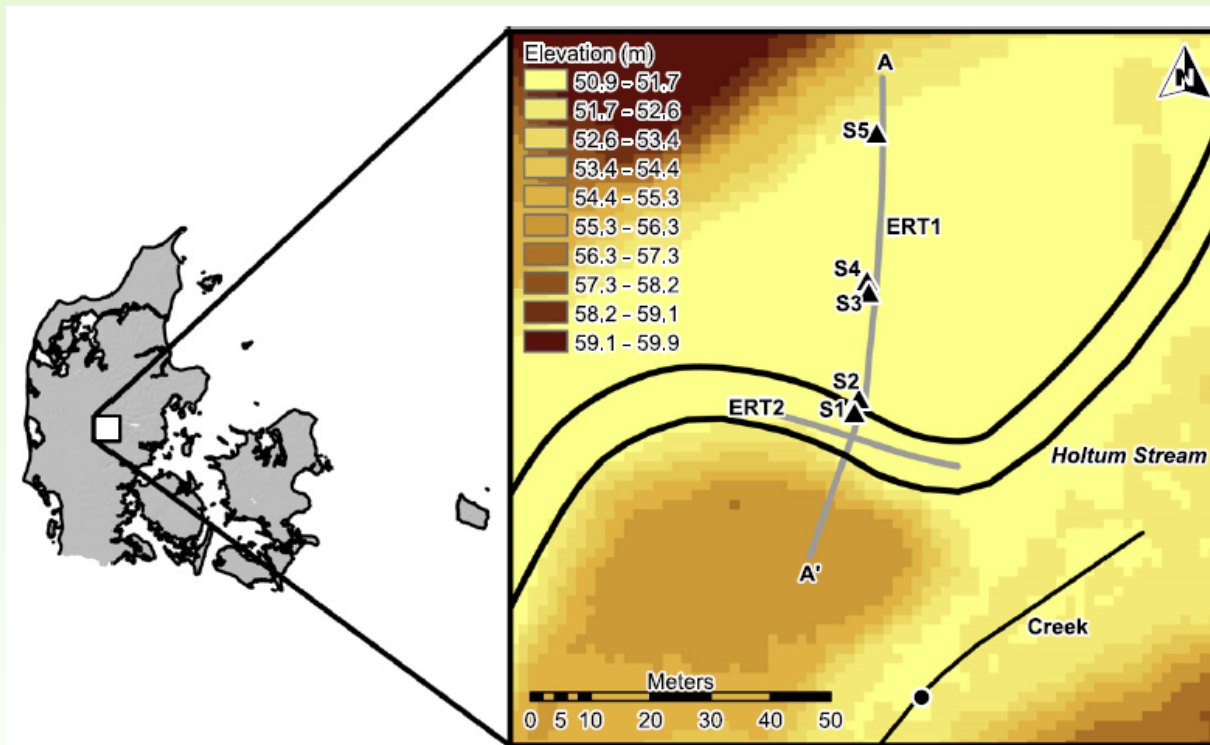


Methods comparison (published in 2020)



Cremeans and Devlin, 2017,
J. Contam. Hydrol.

Field example: A reminder that broader-scale heterogeneity is still important to hyporheic exchange, in this case to nutrient transport to the stream

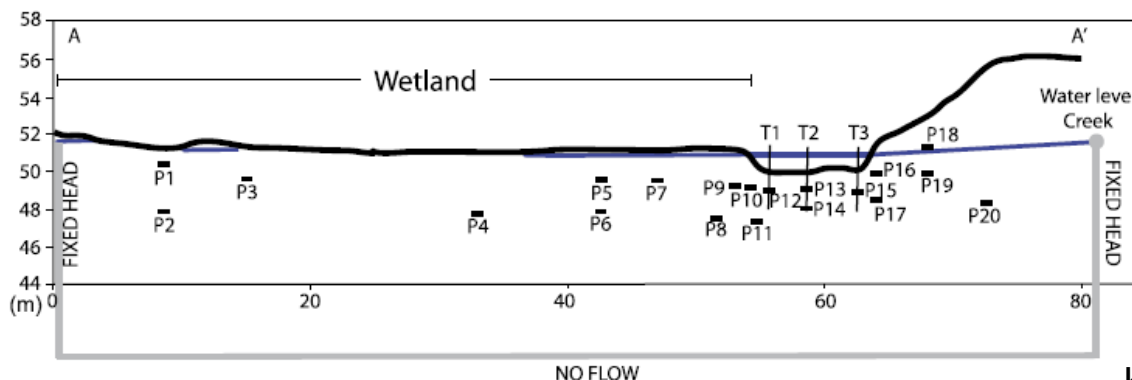


Combined ERT and slug tests from wells to map K , measured seepage in-situ.

Modeled flow to determine flowpaths and travel times

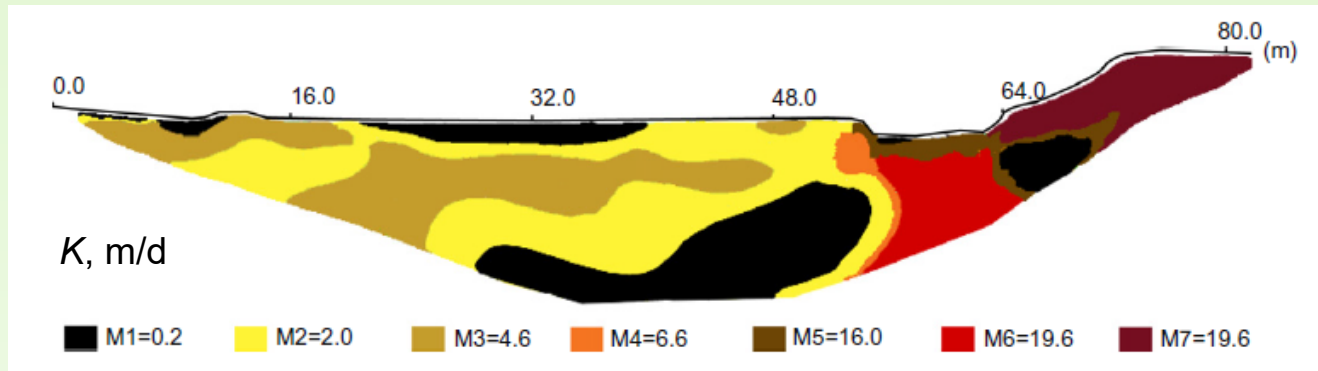
Used CFC and chemistry to confirm the model

A nice example of iteration; field data to understand the local system and constrain a model, then modeling to understand flowpaths and travel times, then data to confirm the modeling.

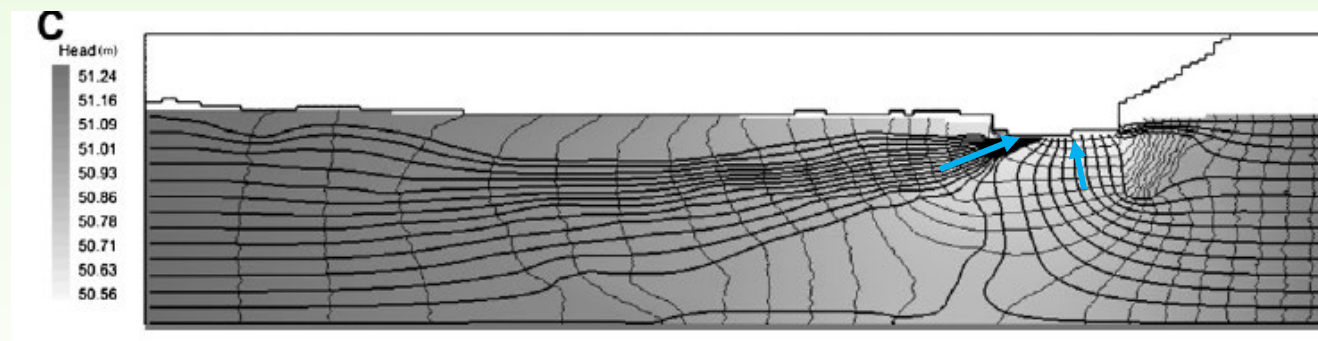


ERT is electrical resistivity tomography. CFC is chlorofluorocarbon, a class of compounds of carbon, hydrogen, chlorine, and fluorine, gases typically used in refrigerants and aerosol propellants. They are harmful to the ozone layer in the earth's atmosphere and were banned in 2000. In some settings, they serve as a good tracer.

Once again, “geology trumps topography”



Results based on field data



Results based on numerical model

Instead of oxic GW discharging on the right side, most discharge was old, nutrient-laden, and was coming from the left side.

Knowing distribution of K was very important to this study. A homogeneous model would not have predicted the distribution of flow and the mix of oxic versus anoxic groundwater discharge to the stream. This affects the proportions of discharging groundwater subject to denitrification. Also, hydraulic head alone was not able to constrain the model during model calibration. Seepage measurements and geochemistry (including CFC age-dating) provided additional information that helped with model calibration.

K distribution changes the distribution of GW discharge to the stream, which affects flowpath length, degree of oxygenation of groundwater, and nitrogen loading to the stream

Here are a few examples of hyporheic processes research priorities (according to UK Environmental Agency)

Hyporheic exchange continues to be a very active area of research

- Develop conceptual models of lowland hyporheic zones
- Develop methods for rapid 3-D in-situ monitoring of ground-water flow
- Determine the main controls on hyporheic exchange (in particular, for lowland rivers)
- Do more research on colmation processes
- Determine if GW and hyporheic water actually mix or if the interface simply shifts
- Does geochemical attenuation capacity vary with sediment deposition?
- What is the nature of immature organic carbon in hyporheic sediments and how does OC affect sorption?
- Is there a hyporheic zone in regulated lowland rivers or do the clay-rich sediment prevent hyporheic exchange?
- What is the potential for attenuation of agricultural nutrients (N and P), industrial pollutants, or heavy metals in hyporheic zones?
- What are the controls on phosphorus cycling in hyporheic zones?
- Review and contrast headwaters hyporheos with assemblages found in lowland rivers
- What is the effect of bioturbation on hydraulic conductivity in hyporheic sediments?
- How does hyporheic-zone hydrology relate to microbial processes and biodiversity?
- How do hyporheic-zone sediments relate to microbial processes and biodiversity?
- Do measures of ecological quality based solely on benthic fauna adequately describe the health of lotic systems?
- Up-scale hyporheic studies from reach scale to catchment scale

References cited

- Alexander, M. D., and D. Caissie (2003), Variability and comparison of hyporheic water temperatures and seepage fluxes in a small Atlantic salmon stream, *Ground Water*, 41(1), 72-82.
- Baxter, C. V., and F. R. Hauer (2000), Geomorphology, hyporheic exchange, and selection of spawning habitat by bull trout (*Salvelinus confluentus*), *Canadian Journal of Fisheries and Aquatic Science*, 57, 1470-1481.
- Blaschke, A. P., K.-H. Steiner, R. Schmalfuss, D. Gutknecht, and D. Sengschmitt (2003), Clogging processes in hyporheic interstices of an impounded river, the Danube at Vienna, Austria, *International Review of Hydrobiology*, 88, 397-413.
- Boano, F., Harvey, J.W., Marion, A., Packman, A.I., Revelli, R., Ridolfi, L., and Wörman, A., 2014, Hyporheic flow and transport processes: Mechanisms, models, and biogeochemical implications: Reviews of Geophysics, v. 52, no. 4, p. 603-679.
- Brown, L. R., and T. Ford (2002), Effects of flow on the fish communities of a regulated California river: Implications for managing native fishes, *River Research and Applications*, 18, 331-342.
- Broecker, T., Sobhi Gollo, V., Fox, A., Lewandowski, J., Nützmänn, G., Aron, S. and Hinkelmann, R. (2021), High-Resolution Integrated Transport Model for Studying Surface Water–Groundwater Interaction. *Groundwater*, 59: 488-502. <https://doi.org/10.1111/gwat.13071>
- Brunke, M., and T. Gonser (1997), The ecological significance of exchange processes between rivers and groundwater, *Freshwater Biology*, 37(1), 1-33.
- Burkholder, B. K., G. E. Grant, R. Haggerty, T. Khangaonkar, and P. J. Wampler (2008), Influence of hyporheic flow and geomorphology on temperature of a large, gravel-bed river, Clackamas River, Oregon, USA, *Hydrological Processes*, 22, 941-953.
- Cardenas, M.B., 2009, Stream-aquifer interactions and hyporheic exchange in gaining and losing sinuous streams: *Water Resources Research*, v. 45, p. W06429, doi:10.1029/2008WR007651.
- Cardenas, M. B., and J. L. Wilson (2006), The influence of ambient groundwater discharge on exchange zones induced by current–bedform interactions, *Journal of Hydrology*, 331, 103-109.
- Cardenas, M. B., and J. L. Wilson (2007), Exchange across a sediment-water interface with ambient groundwater discharge, *Journal of Hydrology*, 346(3-4), 69-80.
- Cardenas, M. B., J. L. Wilson, and V. A. Zlotnik (2004), Impact of heterogeneity, bed forms, and stream curvature on subchannel hyporheic exchange, *Water Resour. Res.*, 40(8), W08307, doi:10.1029/2004WR003008.
- Creameans, M.M., and Devlin, J.F., 2017, Validation of a new device to quantify groundwater-surface water exchange: *Journal of Contaminant Hydrology*, v. 206, p. 75-80.
- Duff, J. H., S. P. Hendricks, A. P. Jackman, and F. J. Triska (2002), The effect of *Elodea canadensis* beds on porewater chemistry, microbial respiration, and nutrient retention in the Shingobee River, Minnesota, North America, *Verhandlungen Internationale Vereinigung für theoretische und angewandte Limnologie*, 28, 1-9.
- Elliott, A. H., and N. H. Brooks (1997), Transfer of nonsorbing solutes to a streambed with bed forms: laboratory experiments, *Water Resour. Res.*, 33(1), 137-151.
- Garrett, J. W., D. H. Bennett, F. O. Frost, and R. F. Thurow (1998), Enhanced Incubation Success for Kokanee Spawning in Groundwater Upwelling Sites in a Small Idaho Stream, *North American Journal of Fisheries Management*, 18(4), 925-930.
- Gooseff, M. N., J. K. Anderson, S. M. Wondzell, J. LaNier, and R. Haggerty (2006), A modelling study of hyporheic exchange pattern and the sequence, size, and spacing of stream bedforms in mountain stream networks, Oregon, USA, *Hydrological Processes*, 20, 2443-2457.
- Grannemann, N. G., and J. M. Sharp Jr. (1979), Alluvial hydrogeology of the lower Missouri River valley, *Journal of Hydrology*, 40, 85-99.
- Harvey, J. W., and K. E. Bencala (1993), The effect of streambed topography on surface-subsurface water exchange in mountain catchments, *Water Resour. Res.*, 29(1), 89-98.
- Jin, G., Chen, Y., Tang, H., Zhang, P., Li, L., and Barry, D.A., 2019, Interplay of hyporheic exchange and fine particle deposition in a riverbed: *Advances in Water Resources*, v. 128, p. 145-157.
- Karan, S., P. Engesgaard, M. C. Looms, T. Laier and J. Kazmierczak. 2013. Groundwater flow and mixing in a wetland–stream system: Field study and numerical modeling. *Journal of Hydrology*, 488:73-83.
- Kaser, D.H., Binley, A., Heathwaite, A.L., and Krause, S., 2009, Spatio-temporal variations of hyporheic flow in a riffle-step-pool sequence: *Hydrological Processes*, v. 23, p. 2138-2149.
- Käser, D., 2010, A new habitat of subsurface waters: the hyporheic biotope (translation of Orghidan's 1959 paper): *Fundamental and Applied Limnology*, v. 176, p. 291-302.
- Lautz, L. K., and D. I. Siegel (2006), Modeling surface and ground water mixing in the hyporheic zone using MODFLOW and MT3D, *Advances in Water Resources*, 29, 1618-1633.

References cited - continued

- Lewandowski, J., Arnon, S., Banks, E., Batelaan, O., Betterle, A., Broecker, T., Coll, C., Drummond, J.D., Gaona Garcia, J., Galloway, J., Gomez-Velez, J., Grabowski, R.C., Herzog, S.P., Hinkelmann, R., Höhne, A., Hollender, J., Horn, M.A., Jaeger, A., Krause, S., Löchner Prats, A., Magliozzi, C., Meinikmann, K., Mojarrad, B.B., Mueller, B.M., Peralta-Maraver, I., Popp, A.L., Posselt, M., Putschew, A., Radke, M., Raza, M., Riml, J., Robertson, A., Rutere, C., Schaper, J.L., Schirmer, M., Schulz, H., Shanafield, M., Singh, T., Ward, A.S., Wolke, P., Wörman, A., and Wu, L., 2020, Is the hyporheic zone relevant beyond the scientific community?: *Water*, v. 11, no. 11, p. 2230.
- Lindgren, R. J., and M. K. Landon (2000), Effects of ground-water withdrawals on the Rock River and associated valley aquifer, eastern Rock County, Minnesota, *Water-Resources Investigations Report*, 103 p pp, U.S. Geological Survey, Mounds View, MN.
- Markstrom, S. L., R. G. Niswonger, R. S. Regan, D. E. Prudic, and P. M. Barlow (2008), GSFLOW—Coupled ground-water and surface-water flow model based on the integration of the precipitation-runoff modeling system (PRMS) and the modular ground-water flow model (MODFLOW-2005), *Techniques and Methods*, 240 pp, U.S. Geological Survey.
- Moir, H. J., C. Soulsby, and A. F. Youngson (2002), Hydraulic and sedimentary controls on the availability and use of Atlantic salmon (*Salmo salar*) spawning habitat in the River Dee system, north-east Scotland, *Geomorphology*, 45, 291-308.
- Movahedi, N., Dehghani, A.A., Schmidt, C., Trauth, N., Pasternack, G.B., Stewardson, M.J., and Meftah Halghi, M., 2021, Hyporheic exchanges due to channel bed and width undulations: *Advances in Water Resources*, v. 149, p. 103857.
- Orghidan, T., 2010, A new habitat of subsurface waters: the hyporheic biotope: *Fundamental and Applied Limnology / Archiv für Hydrobiologie*, v. 176, no. 4, p. 291-302. (Translation by Daniel Käser of paper originally published in German in 1959)
- Packman, A. I., and J. S. MacKay (2003), Interplay of stream-subsurface exchange, clay particle deposition, and streambed evolution, *Water Resour. Res.*, 39(4), 1097, doi:10.1029/2002WR001432.
- Pitlick, J., and M. M. Van Steeter (1998), Geomorphology and endangered fish habitats of the upper Colorado River 2. Linking sediment transport to habitat maintenance, *Water Resour. Res.*, 34(2), 303-316.
- Ridgway, M. S., and P. J. Blanchfield (1998), Brook trout spawning areas in lakes, *Ecology of Freshwater Fish*, 7(140-145).
- Rosenberry, D.O., and Morin, R.H., 2004, Use of an electromagnetic seepage meter to investigate temporal variability in lake seepage: *Ground Water*, v. 42, no. 1, p. 68-77.
- Rosenberry, D.O., and Pitlick, J., 2009, Local-scale spatial and temporal variability of seepage in a shallow gravel-bed river: *Hydrological Processes*, v. 23, p. 3306-3318.
- Ryan, R. J., and M. C. Boufadel (2006), Influence of streambed hydraulic conductivity on solute exchange with the hyporheic zone, *Environmental Geology*, 51, 203-210.
- Salehin, M., A. I. Packman, and M. Paradis (2004), Hyporheic exchange with heterogeneous streambeds: laboratory experiments and modeling, *Water Resour. Res.*, 40(11), W11504, doi:10.1029/2003WR002567.
- Smith, J. W. N. (2005), Groundwater–surface water interactions in the hyporheic zone, *Science Report*, 65 pp, Environment Agency, Almondsbury.
- Song, J., X. Chen, C. Cheng, S. Summerside, and F. Wen (2007), Effects of hyporheic processes on streambed vertical hydraulic conductivity in three rivers of Nebraska, *Geophys. Res. Lett.*, 34, L07409, doi:10.1029/2007GL029254.
- Soulsby, C., A. F. Youngson, H. J. Moir, and I. A. Malcolm (2001), Fine sediment influence on salmonid spawning habitat in a lowland agricultural stream: a preliminary assessment, *The Science of the Total Environment*, 265, 295-307.
- Stanford, J. A. (1998), Rivers in the landscape: introduction to the special issue on riparian and groundwater ecology, *Freshwater Biology*, 40, 402-406.
- Thibodeaux, L. J., and J. D. Boyle (1987), Bedform-generated convective transport in bottom sediment, *Nature*, 325, 341-343.
- Vaux, W. G. (1968), Intragravel flow and interchange of water in a streambed, *Fishery Bulletin*, 66(3), 479-489.
- Veličković, B. (2005), Colmation as one of the processes in interaction between the groundwater and surface water, *Facta Universitatis Series: Architecture and Civil Engineering*, 2(2), 165-172.
- Woessner, W. W. (2000), Stream and fluvial plain ground water interactions: rescaling hydrogeologic thought, *Ground Water*, 38(3), 423-429.